COMPARISON OF FRACTIONATION METHODS FOR SOIL ORGANIC MATTER
14C ANALYSIS

SUSAN E. TRUMBORE and SHUHUI ZHENG

Department of Earth System Science, University of California, Irvine, CA 92697-3100 USA

ABSTRACT. 14C measurements provide a useful test for determining the degree to which chemical and physical fractionation of soil organic matter (SOM) are successful in separating labile and refractory organic matter components. Results from AMS measurements of fractionated SOM made as part of several projects are summarized here, together with suggestions for standardization of fractionation procedures. Although no single fractionation method will unequivocally separate SOM into components cycling on annual, decadal and millennial time scales, a combination of physical (density separation or sieving) and chemical separation methods (combined acid and base hydrolysis) provides useful constraints for models of soil carbon dynamics in several soil types.

INTRODUCTION

Models describing the dynamics of accumulation and turnover of organic carbon generally recognize components of soil organic matter (SOM) that turn over on annual (active), decadal (slow) and centennial to millennial (passive) time scales (Jenkinson and Raynor 1977; O'Brien and Stout 1978; Parton et al. 1987; Jenkinson, Adams and Wild 1991; Potter et al. 1993; Townsend, Vitousek and Trumbore 1995; Schimel et al. 1994). Although these concepts have proven useful in explaining the magnitude and timing of changes in SOM following a perturbation such as land clearing for cultivation (Parton et al. 1987; Cambardella and Elliott 1993; Davidson and Ackerman 1993; Schimel et al. 1994), no recognized method now exists for determining how to apportion SOM into compartments that turn over on different time scales. To be useful for dividing SOM into components for ecosystem modeling, a procedure must determine both the pool sizes and turnover rates.

Two methods are presently in use to partition SOM into active, slow and passive pools. The first model the observed increase of 14C during the 30 yr since the end of atmospheric weapons testing, using additional constraints derived from knowledge of carbon inputs to the system (O'Brien and Stout 1978; Harkness, Harrison and Bacon 1986; Balesdent 1987; Scharpenseel et al. 1989; Trumbore, Vogel and Southon 1989; Vitorello et al. 1989; Jenkinson, Adams and Wild 1991; Harrison, Broecker and Bonani 1993; Trumbore 1993; Townsend, Vitousek and Trumbore 1995; Trumbore, Chadwick and Amundson 1996). The second relies on physical and/or chemical fractionation methods to separate SOM into pools that turn over on different time scales (Paul et al. 1964; Campbell et al. 1967; Scharpenseel, Ronzani and Pietig 1968; Martel and Paul 1974; Goh et al. 1976; Goh, Stout and Rafter 1977; Anderson and Paul 1984; Goh, Stout and O'Brien 1984; Scharpenseel et al. 1989; Trumbore, Vogel and Southon 1989; Trumbore, Bonani and Wölfli 1990; Scharpenseel and Becker-Heidmann 1992).

Radiocarbon measurements of pre-bomb SOM fractions are interpreted as indicative of their turnover rates, though this can cause serious overestimation of the turnover time in soils collected before 1963 (Balesdent 1987; Trumbore 1993; Trumbore and Druffel 1995). A test of the efficiency of any fractionation method at isolating a relatively homogeneous (in terms of turnover) organic matter pool is a comparison of the observed 14C increase since atmospheric weapons testing (Goh, Stout and O'Brien 1984; Balesdent 1987; Trumbore, Vogel and Southon 1989; Trumbore 1993). These tests have shown that, although no fractionation method is completely successful, separation of organic matter by density, followed by hydrolysis in 6N HCl, does leave a residual collection of compounds, which, on average, turn over more slowly than the hydrolyzed portions of SOM (Campbell et al. 1967; Scharpenseel, Ronzani and Pietig 1968; Martel and Paul 1974; Goh et al. 1976;

Here we compare measurements of $^{14}$C made of SOM fractionated using different physical and chemical separation methods. We analyzed both archived (pre-bomb) and contemporary soil samples. The implications of these data for soil carbon cycling are discussed in other publications (see Table 1 for references).

<table>
<thead>
<tr>
<th>Sample</th>
<th>Location</th>
<th>Year collected</th>
<th>Clay content</th>
<th>Soil order/horizon</th>
<th>Vegetation/climate</th>
<th>Parent material</th>
<th>Depth interval (cm)</th>
<th>Ref.*</th>
</tr>
</thead>
<tbody>
<tr>
<td>BS-7</td>
<td>Paragominas, Brazil</td>
<td>1992</td>
<td>&gt;85%</td>
<td>Oxisol/A</td>
<td>Tropical forest</td>
<td>Kaolinitic sediments</td>
<td>30–35</td>
<td>1</td>
</tr>
<tr>
<td>BS-9</td>
<td>Pará, Brazil</td>
<td>1992</td>
<td>&gt;85%</td>
<td>Oxisol/B</td>
<td>Tropical forest</td>
<td>Kaolinitic sediments</td>
<td>100–105</td>
<td>1</td>
</tr>
<tr>
<td>H-13</td>
<td>Lapahoehe, Hawaii</td>
<td>1992</td>
<td>ND</td>
<td>Andisol/A</td>
<td>Tropical pasture</td>
<td>Volcanic ash (15–25 ka)</td>
<td>0–10</td>
<td>2</td>
</tr>
<tr>
<td>H-15</td>
<td>Lapahoehe, Hawaii</td>
<td>1992</td>
<td>ND</td>
<td>Andisol/B</td>
<td>Tropical pasture</td>
<td>Volcanic ash (15–25 ka)</td>
<td>40</td>
<td>2</td>
</tr>
<tr>
<td>NS-11</td>
<td>Shaver Lake, Calif.</td>
<td>1992</td>
<td>7–8%</td>
<td>Ultisol/A</td>
<td>Dry temperate forest</td>
<td>Granodiorite</td>
<td>5–10</td>
<td>3</td>
</tr>
<tr>
<td>NS-13</td>
<td>Shaver Lake, Calif.</td>
<td>1992</td>
<td>7–8%</td>
<td>Ultisol/B</td>
<td>Dry temperate forest</td>
<td>Granodiorite</td>
<td>20–40</td>
<td>3</td>
</tr>
<tr>
<td>OS-9</td>
<td>Shaver Lake, Calif.</td>
<td>1959</td>
<td>7–8%</td>
<td>Ultisol/A</td>
<td>Dry temperate forest</td>
<td>Granodiorite</td>
<td>0–15</td>
<td>3</td>
</tr>
<tr>
<td>OS-10</td>
<td>Shaver Lake, Calif.</td>
<td>1959</td>
<td>7–8%</td>
<td>Ultisol/B</td>
<td>Dry temperate forest</td>
<td>Granodiorite</td>
<td>15–58</td>
<td>3</td>
</tr>
<tr>
<td>A1-52</td>
<td>Michigan</td>
<td>1992</td>
<td>&lt;2%</td>
<td>Anspodosol/B</td>
<td>Moist temperate forest</td>
<td>Beach sand, 3 ka</td>
<td>52</td>
<td>4</td>
</tr>
<tr>
<td>B1-56</td>
<td>Michigan</td>
<td>1992</td>
<td>&lt;2%</td>
<td>Anspodosol/B</td>
<td>Moist temperate forest</td>
<td>Beach sand, 10 ka</td>
<td>56</td>
<td>4</td>
</tr>
</tbody>
</table>

* References:
1. Trumbore et al. (1995)
3. Trumbore, Chadwick and Amundson (1996)

**METHODS**

**Summary of Fractionation Methods**

All samples are sieved to remove components >2 mm. Chemical and density separation procedures have been described before (Trumbore, Vogel and Southon 1989; Trumbore 1993), and are summarized in Figure 1. Three fractionation techniques are used: separation by density, chemical extraction using acids and bases, and separation by size.

**Density Separation**

The density separation extracts organic matter of low density (<1.6–2.0 g cm$^{-3}$) by flotation in a heavy liquid. The low-density material consists of plant matter and charcoal (Sollins, Spycher and
Comparison of Fractionation Methods

Topik 1983; Spycher, Sollins and Rose 1983). Dense fractions contain mineral-associated organic matter, and some microbial cell debris (Sollins, Spycher and Topik 1983; Spycher, Sollins and Rose 1983). In general, the low-density material has higher \(^{14}\text{C}\) values and turns over more rapidly than much of the SOM. Charcoal may be isolated by dissolving other low-density material (except pollen) by treatment with strong acids and bases (Hammond et al. 1991; Gillespie et al. 1992a; Gillespie et al. 1992b).

![Fractionation Scheme](image)

---

**Chemical Fractionation**

Operationally defined fractionation procedures for SOM often rely on extraction using acids and bases. More labile components are hypothesized to be soluble, leaving behind more refractory (and therefore more \(^{14}\text{C}\)-depleted) constituents. These chemical extraction procedures will also affect soil mineral content and can cause changes in SOM chemistry and structure. Extraction with Na-pyrophosphate \((\text{Na}_4\text{P}_2\text{O}_7 \cdot 10\text{H}_2\text{O})\) and 6N HCl will dissolved sesquioxide minerals and remove organic matter complexed with iron and aluminum, and heating SOM in strong acid can cause condensation reactions. The data presented here compare the \(^{14}\text{C}\) content of residues after hydrolysis in acid (HCl)
of varying strength, and in base (NaOH-Na-pyrophosphate). Because the extraction of fresh plant material by these techniques leaves a residue (Ertel and Hedges 1984), we chose to perform acid and base hydrolysis on SOM that has been separated previously by density, and to extract only the denser portion of the soil. For organic-rich Andisol samples, we extracted the bulk soil (as nearly all had densities >2.1 g cm$^{-3}$). Soils developed on basic volcanic rock and containing minerals like allophane and imogolite should employ lower density liquids than we used for improved separation of mineral and organic constituents (Sollins, Spycher and Topik 1983; Spycher, Sollins and Rose 1983).

Size Separation

Recent research has tied the dynamics of decadal-cycling SOM to different size fractions (Cambardella and Elliott 1993, 1994), which has led to speculation about the role of soil aggregates in limiting the decomposition of reactive SOM constituents by physically protecting organic matter from soil microbial activity (Oades 1993). In addition, the limited measurements of $^{14}$C in size-fractionated organic matter available from the literature (Anderson and Paul 1984) suggest a relation between the size of soil-associated organic material and stability of associated organic matter. Recently, Mayer (1994) has shown that organic matter content is related to surface area in many soils and sediments. If organic matter is protected by stabilization on clay surfaces, we might expect it to have lower $^{14}$C values than organic matter, which is more available for microbial consumption.

For the size separation, we used the <2-mm sieved samples, without prior density separation. We measured the $^{14}$C content of three size fractions: <2 mm but >63 $\mu$m, <63 $\mu$m but >2 $\mu$m, and <2 $\mu$m. The largest fraction was separated by dry sieving; the fractions > and <2 $\mu$m were separated by settling from a suspension (Folk 1961). This procedure breaks apart soil aggregates.

Carbon and $^{14}$C Measurements

Percent carbon data were determined by CO$_2$ evolution on in-vacuo combustion at 900°C with cupric oxide wire (Buchanan and Corcoran 1959). $^{14}$C analyses were made by accelerator mass spectrometry (AMS) at the Center for AMS, Lawrence Livermore Laboratory, Livermore, California (Southon et al. 1992). Graphite targets for AMS measurement were prepared from the purified CO$_2$ using the sealed-tube, zinc reduction method modified from Vogel (1992). $^{14}$C data are reported as $\Delta^{14}$C, the per mil difference in $^{14}$C/$^{12}$C ratio between the sample and an absolute standard (oxalic acid decay-corrected to 1950) (Stuiver and Polach 1977; Donahue, Linick and Jull 1990). $\delta^{13}$C values used to correct for mass-dependent fractionation effects were assumed to be $-25\%$ for all samples, as all sites had predominantly C$_3$ vegetation.

Nomenclature

We adopt a shorthand for describing the fractionation treatment experienced by a sample. The descriptors, denoted in Figure 1, are defined as follows:

- BS = bulk soil. Due the expense of AMS measurements, we often rely on measurements of the constituent fractions to derive a bulk soil $^{14}$C number by mass balance.
- LD = low-density fraction (< 2.0 to 2.1 g cm$^{-3}$).
- HD = dense fraction (> 2.0 to 2.1 g cm$^{-3}$).

All chemical fractionation procedures are performed on the HD fraction, and the data associated with them refers to the residue after a given chemical separation procedure. A refers to acid hydrolysis at room temperature in 0.5 N HCl; B refers to extraction with 0.1 N NaOH–0.1 N sodium pyrophosphate; C refers to hydrolysis in 6N HCl at 95°C. Often, we show data derived from the residue
of a combination of chemical fractionation procedures: for example, A+B+C refers to the residue after sequential hydrolysis in 0.5 N acid, 0.1 N NaOH and 6 N HCl. When performing these procedures, the residue after each extraction step is rinsed thoroughly in distilled H2O before moving on to the next procedure.

**Samples**

We present here summaries of fractionation for several different kinds of soils. Table 1 summarizes data related to soil classification and site-specific characteristics, such as climate and parent material. References discussing the carbon cycling in these soils, or containing more information about the soils, as well as more detailed discussion of 14C data with regard to soil carbon dynamics at each site, are also listed.

**RESULTS**

Tables 2 and 3 show the carbon and 14C content of the fractionated SOM. To better compare the effects of fractionation methods on the different soils, the results are also plotted in Figures 2 and 3.

**Density Separation**

Low-density components consist primarily of fine root hairs, charcoal (found in most LD fractions, but not quantified separately) and surface detritus mixed into the soil (in Oxisols BS-7, BS-9). Low-density material makes up a large portion of the total soil carbon in A horizons of more coarse-textured soils and the Andisol (which also had the highest total carbon content, Table 2). The 14C content of LD material is close to atmospheric 14CO2 values at the time of collection in both pre- and post-bomb soils, indicating rapid turnover of this fraction.

In general, 14C values of the dense fraction (mineral-associated) carbon decreased with depth in the soil, as has been observed (Scharpenseel et al. 1989). A notable exception is in the Spodosol Bs horizon (where sesquioxide accumulation is occurring), which shows increased 14C values compared to the overlying eluviated horizon. Bomb 14C is present in HD fraction for soils sampled in the 1990s, indicating the presence of at least some fast-cycling carbon associated with soil mineral phases.

**Acid-Base Hydrolysis**

In the heavier-textured soils (excluding the Shaver soil from the Sierra Nevada, California), most of the carbon present is associated with mineral surfaces (HD fraction). Hydrolysis of the HD fraction in relatively mild acids (0.5 N HCl) removes >30% of the mineral-associated carbon for all soils studied (see Fig. 2A). The removal is greatest in Spodic B horizons (the Michigan soils samples, B1-56 and A1-52), where >70% of the organic carbon is solubilized in 0.5 N HCl. The carbon remaining is depleted in 14C compared to the starting dense fraction material to varying degrees (Fig. 3A). Least affected are the Andisol (Hawaii) samples; most affected are the spodic B horizons. Subsequent hydrolysis in base and strong acid removes more carbon and 14C (Figs. 2B and 3A–C). The residue after A+B+C treatments is, in all but two cases, the most depleted in 14C of any treatment. Figure 3 summarizes the difference in 14C content of residual carbon after various treatments, compared to the starting dense material (1:1 line on figures). The A+B treatment (Fig. 3B) removes almost as much carbon and 14C as the A+B+C treatment, while treatments using only acids (A+C; Fig. 3C) do not remove as much C and 14C as treatments including a base hydrolysis step. The Andisols and Oxisols are affected by chemical fractionation less than coarser-textured soils. The difference between unfractionated dense material and the A+B or A+B+C residues is smaller in A horizons of soils than in B horizons. This is more noticeable in previously published profiles (Trumbore, Bonani and Wölfli 1990).
### TABLE 2. Results of Chemical Fractionation Procedures*

<table>
<thead>
<tr>
<th>Sample</th>
<th>C density (gC m(^{-2}) cm(^{-1}))</th>
<th>LD &lt;2.0 g/cc</th>
<th>D (&gt;2 g/cc)</th>
<th>A only</th>
<th>A + B</th>
<th>A + B + C</th>
<th>A + C</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>%C(_{LD})</td>
<td>(\Delta^{14}C)</td>
<td>%C(_g)</td>
<td>(\Delta^{14}C)</td>
<td>%C(_d)</td>
<td>(\Delta^{14}C)</td>
<td>%C(_d)</td>
</tr>
<tr>
<td>BS-7</td>
<td>132</td>
<td>&lt;5%</td>
<td>+170</td>
<td>1.0</td>
<td>-102.4</td>
<td>61.7</td>
<td>-195.0</td>
</tr>
<tr>
<td>BS-9</td>
<td>50</td>
<td>&lt;1%</td>
<td>ND</td>
<td>0.4</td>
<td>-552.0</td>
<td>72.0</td>
<td>-615.0</td>
</tr>
<tr>
<td>H-13</td>
<td>ND</td>
<td>ND</td>
<td>ND</td>
<td>11.0</td>
<td>4.0</td>
<td>46.5</td>
<td>-16.3</td>
</tr>
<tr>
<td>H-15</td>
<td>ND</td>
<td>ND</td>
<td>ND</td>
<td>4.2</td>
<td>-660.8</td>
<td>62.1</td>
<td>-705.9</td>
</tr>
<tr>
<td>NS-11</td>
<td>462</td>
<td>91.7</td>
<td>+109.3</td>
<td>1.5</td>
<td>92.6</td>
<td>12.2</td>
<td>4.9</td>
</tr>
<tr>
<td>NS-13</td>
<td>98</td>
<td>73.5</td>
<td>+169.5</td>
<td>0.4</td>
<td>32.7</td>
<td>40.9</td>
<td>-55.2</td>
</tr>
<tr>
<td>OS-9</td>
<td>383</td>
<td>85.1</td>
<td>-35.8</td>
<td>2.1</td>
<td>-36.5</td>
<td>13.0</td>
<td>-93.9</td>
</tr>
<tr>
<td>OS-10</td>
<td>92</td>
<td>59.8</td>
<td>-30.9</td>
<td>0.4</td>
<td>-123.0</td>
<td>39.4</td>
<td>-145.3</td>
</tr>
<tr>
<td>A1-52</td>
<td>ND</td>
<td>&lt;5%</td>
<td>ND</td>
<td>0.7</td>
<td>-148.7</td>
<td>25.6</td>
<td>-298.6</td>
</tr>
<tr>
<td>B1-56</td>
<td>ND</td>
<td>&lt;5%</td>
<td>ND</td>
<td>1.0</td>
<td>-123.7</td>
<td>12.4</td>
<td>-231.2</td>
</tr>
</tbody>
</table>

*Sample identification given in Table 1. Carbon density (CD; gC m\(^{-2}\) cm\(^{-1}\)) is calculated from bulk density and %C. ND = no data. The first fractionation procedure was separation by density into <2.0 g/cc (LD) and >2.0 g/cc (HD) fraction. %C\(_{LD}\) denotes the percent of total carbon in the LD fraction. The percent of total carbon in HD fraction (not shown) is 100-%C\(_{LD}\). Hawaii (H-13 and H-15) samples were not separated by density. For these soils, the bulk data are listed in the ID column, since most of the soil had density <2.0 g/cc. The starting material for chemical fractionation procedures was the HD fraction, for which gravimetric percent carbon (%C\(_d\)) data are shown.

Chemical procedures are as follows:

A = residue after extraction with 0.5 N HCl at room temperature for 12 h.
B = residue after extraction at room temperature for 24 h with 0.1 N NaOH-0.1 N Na\(_2\)P\(_2\)O\(_7\) · 10H\(_2\)O.
C = residue after extraction with 6 N HCl at 95°C for 12 h.

Combinations of chemical procedures are shown, e.g., results under the heading A+B are for the residual organic matter following treatment with first A, then B. %Cd is the percent of the carbon in the untreated HD fraction that remains after chemical treatment. 100% minus the %Cd gives the percent of total HD carbon extracted. \(^{14}C\) data are reported as \(\Delta^{14}C\) (%), assuming \(^{13}C\) of -25% for all samples; vegetation was exclusively C\(_3\) at all sites. The 1-\(\sigma\) uncertainty of the \(^{14}C\) measurements (based on repeated measurements of secondary standards) is ± 10 to 15%.
Fig. 2. Percent of total carbon in the residues left after acid and base hydrolysis treatments of the dense fraction (A) and after size fractionation of the bulk soil (B).
### TABLE 3. Results of Size Fractionation Procedures*

<table>
<thead>
<tr>
<th>Sample</th>
<th>$2 \text{mm} &gt; S &gt; 63 \mu$</th>
<th>$63 \mu &gt; S &gt; 2 \mu$</th>
<th>$&gt; 2 \mu$</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>%C</td>
<td>%Ct</td>
<td>$\Delta^{14}C$</td>
</tr>
<tr>
<td>BS-7</td>
<td>3.9</td>
<td>10.6</td>
<td>7.6</td>
</tr>
<tr>
<td>BS-9</td>
<td>0.4</td>
<td>33.3</td>
<td>-465.8</td>
</tr>
<tr>
<td>NS-11</td>
<td>11.2</td>
<td>61.3</td>
<td>82.9</td>
</tr>
<tr>
<td>NS-13</td>
<td>0.8</td>
<td>69.5</td>
<td>11.0</td>
</tr>
<tr>
<td>OS-9</td>
<td>2.2</td>
<td>63.5</td>
<td>-54.4</td>
</tr>
<tr>
<td>OS-10</td>
<td>0.9</td>
<td>64.2</td>
<td>-108.1</td>
</tr>
<tr>
<td>A1-52</td>
<td>0.6</td>
<td>87.8</td>
<td>-128.8</td>
</tr>
<tr>
<td>B1-56</td>
<td>0.6</td>
<td>93.1</td>
<td>-113.3</td>
</tr>
</tbody>
</table>

*%C is gravimetric carbon content, %Ct is the percent of total soil carbon in the size fraction and $^{14}C$ data are reported as $\Delta^{14}C$ (%), assuming $\delta^{13}C$ values of $-25\%$ for all samples.

---

Fig. 3. $\Delta^{14}C$ of the dense fraction (HD) compared to that measured in the residue after various chemical treatments. The 1:1 line (which signifies no difference in $^{14}C$ between the residue and the untreated original material) and the $\Delta^{14}C$ values for A+B+C treatments (normally the most $^{14}C$-depleted residues) are shown on each figure.
Size Fractionation

The results of size fractionation on splits of the same soils in which density and chemical procedures were used are reported in Table 3. Most of the carbon in coarser-textured soils was associated with the >63 μ fraction (Fig. 2B), whereas in more clay-rich, heavy-textured soils (Oxisols), most of the organic carbon is associated with the fine silt and clay fractions (<63 μ). Size separations were not performed on the Andisol (Hawaii) soils; these are dominated by fine-grained material. Size-fractionation of soils did not affect 14C values as much as chemical fractionation (Fig. 3D). We did not see any consistent tendency for clay-sized particles to have Δ14C values much greater or less than other size fractions, although clay fractions both enriched and depleted in 14C were observed.

SUGGESTIONS FOR A STANDARD FRACTIONATION SCHEME

The usefulness of different fractionation methods depends in part on the type of soil, in particular, on soil texture (clay content). Separation by density will concentrate organic matter that turns over rapidly (decades and shorter time scales) in the low-density fraction. In coarse-textured soils with granitic parent material, such as the Shaver soil (Sierra Nevada), low-density organic matter can make up the majority of the total soil carbon in the A horizon. The remaining HD material shows significant 14C increases in A horizons due to incorporation of bomb 14C. Chemical treatment, especially hydrolysis in acid and base, removes relatively 14C-enriched components of the dense, mineral-associated organic matter. Chemical treatment is most effective in coarse-textured soils, in part because more of the total carbon is removed in the hydrolysis procedure. Chemical extraction with acids and bases did not change the 14C content of Andisol organic matter, even though up to 90–95% of the original carbon was removed.

Large differences in 14C content were not observed between different size fractions. Most of the carbon in coarse-textured soils is in the larger size fractions (>63 μ), and probably is mostly low-density vascular plant material, although 14C values indicate that this organic matter is diluted by other (mineral associated) constituents.

We suggest the most informative fractionation procedure to be used in separating SOM into faster and slower cycling components combines density and chemical separation procedures. Low-density organic material (except for charcoal) turns over on timescales of decades and less (Trumbore 1993; Trumbore, Chadwick and Amundson 1996). Extraction of dense material with acid and base will leave a residue that can provide a minimum estimate of the size and 14C content of a passive pool in 14C modeling. This fractionation will be most useful in soils of medium-to-coarse texture, with granitic parent material.

ACKNOWLEDGMENTS

This research was undertaken as part of projects funded by NASA’s Terrestrial Ecology and Earth Science Programs and a grant from the A. W. Mellon Foundation. We are extremely grateful to John Souther, Michaele Kashgarian, Jay Davis, and others at the Center for Accelerator Mass Spectrometry of Lawrence Livermore National Laboratory for continued rapid throughput of 14C AMS samples.

REFERENCES


Barrett, L. R. and Schaetzl, R. J. 1992 An examination of podzolization near Lake Michigan using chronofunc-


---


Sollins, P., Spycher, G. and Topik, C. 1983 Processes of
Comparison of Fractionation Methods


