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Zhang, Y.

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Single-well experimental design for studying residual trapping of supercritical carbon dioxide

Yingqi Zhang¹, Barry Freifeld¹, Stefan Finsterle¹, Martin Leahy²,³, Jonathan Ennis-King²,³, Lincoln Paterson²,³, Tess Dance²,³

¹Lawrence Berkeley National Laboratory, Berkeley, CA, USA
²CSIRO Petroleum, Clayton, Victoria, Australia
³Cooperative Research Centre for Greenhouse Gas Technologies, Australia

Abstract
The objective of our research is to design a single-well injection withdrawal test to evaluate residual phase trapping at potential CO₂ geological storage sites. Given the significant depths targeted for CO₂ storage and the resulting high costs associated with drilling to those depths, it is attractive to develop a single well test that can provide data to assess reservoir properties and reduce uncertainties in the appraisal phase of site investigation. The main challenges in a single-well test design include (1) difficulty in quantifying the amount of CO₂ that has dissolved into brine or migrated away from the borehole; (2) non-uniqueness and uncertainty in the estimate of the residual gas saturation ($S_{gr}$) due to correlations among various parameters; and (3) the potential biased $S_{gr}$ estimate due to unaccounted heterogeneity of the geological medium. To address each of these challenges, we propose (1) to use a physical-based model to simulation test sequence and inverse modeling to analyze data information content and to quantify uncertainty; (2) to jointly use multiple data types generated from different kinds of tests to constrain the $S_{gr}$ estimate; and (3) to reduce the sensitivity of the designed tests to geological heterogeneity by conducting the same test sequence in both a water-saturated system and a system with residual gas saturation. To perform the design calculation, we build a synthetic model and conduct a formal analysis for sensitivity and uncertain quantification. Both parametric uncertainty and geological uncertainty are considered in the analysis. Results show (1) uncertainty in the estimation of $S_{gr}$ can be reduced by jointly using multiple data types and repeated tests; and (2) geological uncertainty is essential and needs to be accounted for in the estimation of $S_{gr}$ and its uncertainty. The proposed methodology is applied to the design of a CO₂ injection test at CO2CRC’s Otway Project Site, Victoria, Australia.

1. Introduction and Objective
The geologic sequestration of anthropogenic greenhouse gases to mitigate climate change is receiving increasing attention as a means to reduce atmospheric emissions and the related impacts as a result of continued use of fossil fuels. The ability of a host formation to effectively trap CO₂ determines the suitability of a proposed site for long-term CO₂ sequestration. Four trapping mechanisms have been identified (IPCC, 2005): structural trapping, residual phase trapping, solubility trapping and mineralization trapping. This study focuses on residual phase trapping, i.e., the immobilization of individual bubbles or relatively small blobs of the CO₂-rich phase. The CO₂ bubbles are either trapped by capillary forces or are stuck in local trapping structures or dead-end portions of the pore space, preventing further CO₂ migration in response to pressure gradients or buoyancy forces. (CO₂ saturation can be reduced below the residual value by processes other than viscous flow, e.g., by compression or dissolution.) A parameter referred to as residual gas saturation ($S_{gr}$) is used to characterize the tendency of a geologic formation to trap some of the non-wetting phase in its pore space. The residual gas saturation is a property of the interaction between the porous medium and the fluids, mostly reflecting the size and shape of its pores and their connectivity. However, residual gas saturation is not a static parameter; it depends on the sequence of hysteretic drainage and imbibition processes, i.e., it is history-dependent, with different values at each point in the storage formation as the fluid saturation changes during CO₂ injection and redistribution. Only its maximum value $S_{gr\text{max}}$ (associated with the primary imbibition curve) can be considered as a formation parameter independent of the dynamic system.
state. The determination of the maximum residual gas saturation of a formation is both challenging and crucial as it will determine the ultimate capacity of a reservoir to residually trap CO₂.

To achieve the pressure and temperature conditions required for supercritical CO₂ storage, the depth of the target formation must have pressures of 72 bar or greater, with a temperature above 31 °C. In most sedimentary basins these conditions occur at depths greater than 800 m below ground surface. The high cost associated with drilling to such depths may limit the number of wells available for site characterization. Consequently, we believe there is significant value in developing a single-well appraisal methodology that can determine parameters critical for CO₂ sequestration projects prior to full scale injection with acceptable accuracy.

The objective of this study is to propose a suite of single-well tests for quantifying the maximum residual gas saturation in brine aquifers. Although the maximum residual gas saturation is heterogeneous, an effective maximum residual gas saturation on the support scale of the test is the parameter more relevant to estimating CO₂ storage capacity, and that is the parameter of interest in this work. The main difficulties in the test design for estimating Sₘₘ are: (1) due to the complexity of multiphase flow processes, there is no standard well test available for determination of Sₘₘ. This requires us to use a process-based data analysis approach, i.e., an inversion analysis that is based on a process model; (2) correlations among various uncertain parameters lead to non-unique estimates of Sₘₘ; e.g., an observed high pressure increase during CO₂ injection may be caused by either a high Sₘₘ or a low formation permeability; therefore, the pressure signal alone does not give us enough information to infer Sₘₘ. This requires us to jointly use multiple types of data that are complementary to each other; (3) heterogeneity of aquifer properties and related uncertainty in fluid distribution after injection may lead to a biased estimate of Sₘₘ. This requires us to design a test that is relatively insensitive to geological heterogeneity. Our proposed test sequence is designed to address the above challenges. We will (1) demonstrate how our approach arrives at a relatively reliable estimate of Sₘₘ through a synthetic example, and (2) demonstrate the approach by applying it for the design of the pilot injection at CO2RC’s Otway Project Site, Victoria, Australia. The proposed method is believed to be generally applicable to other CO₂ injection projects.

2. Design approach

Single-hole testing, while being standard practice in groundwater hydrology, faces considerable challenges when considered for the characterization of multi-phase flow systems. Conventional single-well tests exploit the fact that the pressure perturbation (i.e., the signal analyzed by analytical or numerical well-test analysis methods) propagates through the geologic formation much farther and faster than the injected fluid itself. However, for the characterization of a multiphase system, or (as in this case) the determination of two-phase flow parameters of a formation that is currently under single-phase liquid conditions, it is necessary to create two-phase conditions, where the two-phase region around the wellbore is substantially smaller than the single-phase region that can be interrogated by the pressure perturbation in standard single-phase well testing. Furthermore, highly nonlinear multiphase processes and related phase transitions need to be captured for data analysis.

As a starting point for our test design, we consider a test sequence frequently used for oil reservoir characterization by (1) injecting CO₂, (2) shutting-in the well to allow the injected CO₂ to migrate and dissolve into the oil phase, and (3) producing the mobilized mixture back (Mohammed-Singh et al., 2006; Monger et al., 1991). This “Huff-n-Puff” methodology has been used as part of a CO₂ sequestration demonstration at the Loudin Oil Field (Berger et al., 2009; Frailey and Finley, 2008) as part of an enhanced oil recovery (EOR) operation. If we apply the same procedure to a system that contains only brine (without oil), the amount of CO₂ produced back can be used to approximate the amount of non-trapped CO₂. In other words, the difference between the injected CO₂ and produced CO₂ is considered to be trapped, which in turn may allow determination of the residual gas saturation. This approach has a number of shortcomings. Foremost, the amount of CO₂ not recovered is not necessarily trapped, but may not have been produced because CO₂ either has dissolved into the brine or has migrated away from the injection zone due to buoyancy effects and escaped from the well’s capture zone during production. Therefore, this method tends to overestimate long-term residual trapping. Even if the total volume of trapped CO₂ were properly estimated, its spatial distribution in the formation remains unknown. Consequently, this method yields great uncertainty in the estimate of the residual CO₂ saturation, which is the parameter of interest. Assumptions about the efficiency of brine displacement, and the geometry of the trapped CO₂ plume after pull-back may be too simplistic for an accurate estimation of Sₘₘ. In summary, simply calculating the mass balance, i.e., the difference between the amount of gas injected and recovered, may yield an upper estimate of trapped CO₂ in a specific single-hole test, but does not allow the unique and accurate determination of residual gas saturation as a parameter that can be used for predictive simulations of CO₂ trapping during sequestration operations.

Another potentially useful test is a partitioning tracer test run in an injection-withdrawal (push-pull) configuration from a single well. Partitioning tracer tests that are done between two wells are referred to as partitioning interwell tracer test (PITT). PITTs have been developed to detect crude oil (Cook, 1971; Dean, 1971; Tang and Harker, 1991), measure water saturation (Imhoff et al., 2003; Li and Imhoff, 2005), and characterize and quantify nonaqueous phase liquid (e.g., NAPL) saturation (Deeds et al., 1999; Jin et al., 1995; Mariner et al., 1999; Nelson et al., 1999; Whitley et al., 1999; Young et al., 1999). In PITT, tracers are injected upstream and the recovered tracer breakthrough curves (BTCs) at the downstream extraction well are used to analyze tracer retardation, which is used to infer phase saturation. Single-well push-pull partitioning tracer tests have also been conducted to quantify properties for a wide range of aquifer physical, biological, and chemical characteristics, such as estimating solute retardation factors (Schroth et al., 2001); quantifying in situ rates of CH₄ oxidation in soils (Gomez et al., 2008; Gonzalez-Gil et al., 2007; Urmann et al., 2008); and detecting and quantifying the presence of NAPL (Davis et al., 2002, 2005; Istok et al., 2002; Schroth et al., 2001). A single-well partitioning push-pull tracer test includes (1) injecting water containing partitioning and conservative tracers into a well, (2) continuing to inject water to push the injected tracer away and allow tracer to partition into the nonaqueous phase liquid, which is assumed to be immobile, and (3) producing fluid back and measuring tracer concentrations. Because the injected tracer partially partitions into the nonaqueous phase liquid, recovered BTCs of tracers with higher nonaqueous phase liquid solubility are retarded with respect to more conservative tracers that preferably remain in the aqueous phase. Multiple tracers with different partitioning coefficients result in different BTCs that can be analyzed for the estimation of average immobile saturation of nonaqueous phase liquid. Tomich et al. (1973) used a reactive tracer (ethyl acetate) that is soluble in both the water and oil phases to hydrolyze and form ethanol, a second tracer, that is only soluble in water, and used the difference between the two tracers in a push-pull test to calculate the retardation factor and estimate the residual oil saturation. The recovered BTCs were analyzed by plotting the normalized concentration (concentration of extracted fluid divided by the initial injected solution concentration) versus the ratio between the extracted volume and injected volume. Schroth et al. (2001)
used numerical simulations to study the effects of retardation factors and (linear or nonlinear) sorption processes on the recovered BTCs in a homogeneous system. They concluded that the apparent dispersion increases with increased retardation factor. Davis et al. (2002, 2005) used radon that occurs naturally in groundwater as the partitioning tracer, along with injected water that contains a conservative tracer. Istok et al. (2002) analyzed the recovered BTCs of a few tracers from single-well push–pull tests with and without NAPL, to detect the presence of NAPL under both laboratory and field conditions. Although their study demonstrated that single-well partitioning tracer test can be used to detect NAPL, they concluded that additional research is needed to verify the ability of these tests to quantify NAPL saturation. All these studies (Davis et al., 2002, 2005; Istok et al., 2002; Nelson et al., 1999; Schroth et al., 2001) have pointed out that the presence of porous-medium heterogeneity and a variable distribution of immobile phase saturation can lead to reduced accuracy of the partitioning tracer test, and their impact on BTCs needs to be investigated.

Based on the results obtained with NAPLs, it can be expected that a single-well push–pull test also can be applied to quantify residual CO2 saturation in the field. However, estimation of residual gas saturation from tracer retardation is challenging because tracer properties (partitioning coefficient, diffusion coefficient) are uncertain given the high pressure and elevated temperature conditions expected at CO2 injection sites and the measurement of tracer concentrations carries experimental errors. Moreover, the inherent heterogeneity of the geological formation and the assumption of “instant equilibrium partitioning” used to analyze tracer data leads to uncertain and non-unique estimates of residual gas saturation.

The test design proposed below aims at overcoming these limitations. Moreover, the design should ensure that sufficiently sensitive data are collected that contain the information needed to enable inference of the parameters of interest with acceptably low estimation uncertainty. This is achieved through (1) a forward numerical model that is capable of simulating non-isothermal, multicomponent, multiphase flow processes to simulate the test sequence; and (2) an inverse modeling analysis to optimize the test design: to determine information content of data sets through systematic sensitivity analyses and to examine estimation uncertainty using synthetic data inversions. Two principles are followed in the design: (1) perform a test sequence that yields multiple types of complementary data to constrain the estimate of Sr_gr; and (2) remove or reduce the bias caused by the heterogeneity of the storage formation by repeating the same test under different saturation conditions.

We propose to use a combination of three tests performed before and after CO2 injection: a hydraulic test, a thermal test, and a push–pull partitioning tracer test. The basic idea is to (1) perform this series of hydraulic, thermal and tracer tests under fully water-saturated conditions (referred to as reference tests; Stage 1); (2) create a system at residual CO2 saturation (Stage 2); (3) repeat the sequence of hydraulic, thermal and tracer tests in the formation containing residual CO2 (referred to as characterization tests; Stage 3); and (4) analyze data from all three stages and estimate Sr_gr. Performing reference tests provides information for single-phase parameters. As a result, the reduced uncertainty in these parameters will help to constrain two-phase parameters. The individual tests can be described as follows:

A. Hydraulic test: In the hydraulic test, water is injected into the formation (which may be fully liquid saturated or contain gas), and pressure transients are recorded and utilized to infer the amount of gas trapped in the formation. Pressure data are sensitive to residual gas saturation because pressure change depends on relative permeability, which depends on residual gas saturation. When water is injected into (a) a fully water-saturated formation; (b) a formation at residual gas saturation with a low Sr_gr; and (c) a formation at residual gas saturation with a high Sr_gr, the liquid relative permeability is highest in (a) and lowest in (c), which means the pressure pulse in case (a) can move away from the vicinity of the well much faster than in cases (b) and (c). Therefore, the pressure increase in case (c) is the highest, even though the effect is somewhat countered by the increase in storativity as gas saturation is increased. Inherent heterogeneity effects are reduced by comparing the pressure response in a formation at residual gas saturation with the reference test response.

B. Tracer test: Tracer is injected into the formation and pushed further away from the injection well by continuous injection of water, before being pulled back into the borehole, where a breakthrough curve (BTC) is measured. Tracer breakthrough curves are sensitive to residual gas saturation because a portion of tracer will partition into the gas, becoming immobile, and thus will not be produced back. Tracer partitioning between water and gas is governed by Henry’s law, which states that at constant temperature, the amount of a given gas dissolved in a given type and volume of liquid is directly proportional to the partial pressure of that gas in equilibrium with that liquid. This proportionality constant, which is also temperature dependent, is known as Henry’s coefficient. Puess et al. (2005) provides Henry’s coefficients for a list of potential noble gases (such as Ar, Kr, Xe) that can be used as tracers. Once tracer is injected into a formation that contains residual gas, part of the tracer partitions into the gas and becomes immobile (to advective forces). The amount of tracer that remains in the water is mobile and can be produced during the pull-back phase of the test. Then, Sr_gr can be inferred based on the amount of tracer that is recovered. Two comparisons can be made using the tracer breakthrough curves: (1) for the same tracer, we compare the BTCs from Stage 1 and Stage 3, the differences of which can be used to infer Sr_gr; (2) for two tracers with different partitioning coefficients, we can compare their recovery curves from the same test to estimate Sr_gr. Use of both comparisons can help reduce the uncertainty in the estimation of Sr_gr. In this test, we propose to use noble gases as tracers, because (1) they partition between the aqueous and CO2-rich phase, (2) they are chemically inert and non-hazardous, and (3) they can be measured with great precision. We propose to use two noble gases with distinct partitioning coefficients and compare the BTCs obtained from the residual gas field (Stage 3) with the BTCs from the reference test (Stage 1).

C. Thermal test: In the thermal test, the borehole is heated and temperature data are recorded. Temperature data are sensitive to residual gas saturation because temperature change during thermal perturbation depends on thermal diffusivity, which depends on fluid composition, i.e., the amount of CO2 in the system. Supercritical CO2 has a significantly smaller thermal diffusivity than water (e.g., at 65°C and 14.5 MPa, for a sand formation with a porosity of 20%, a specific heat capacity of 800 J/(kg°C), when it is fully water saturated, the effective thermal conductivity is about 2.2 W/(m K) and diffusivity about 8.9 × 10^-7 m^2/s; when it is fully CO2 saturated, the effective thermal conductivity is about 1.4 W/(m K), and its diffusivity is about 6.8 × 10^-7 m^2/s). We can take advantage of the contrast between the thermal diffusivity of a brine-saturated formation and one that contains CO2 by performing a heat transport test in which the temperature response is used to estimate gas saturation. Freifeld et al. (2008) described the basic mechanics of such a test, including the deployment of a fiber–optic distributed temperature sensor (DTS) and a resistance heater, and the analysis of the collected temperature data for estimating formation thermal conductivity in situ. In addition, the application of thermal monitoring to CO2 storage has been demonstrated as part of the CO2SINK experiment (Freifeld et al. 2008).
Fig. 1. Maximum temperature after 2 days of heating at 40 W/m for systems with different gas saturations.

et al., 2009; Giese et al., 2009). Fig. 1 shows the temperature increase as a function of saturation (assuming a uniform distribution of CO₂) after 2 days of heating at 40 W/m of a system that is initially at 65°C. For such a system, a temperature change of 0.1°C corresponds to a saturation difference of 0.02. The temperature increase is directly proportional to gas saturation. One of the benefits of using temperature transients as a proxy measurement to estimate \( S_{gr} \) is the high accuracy and relative simplicity in monitoring thermal response. Given a typical DTS measurement accuracy of ±0.1°C (with a repeatability of 0.03°C), we are able to detect CO₂ saturation changes smaller than 0.05. In addition, the depth of penetration of the thermal front into the formation (1–2 m) is significantly greater than alternative wellbore logging methods (i.e., a pulsed neutron capture tool, which only penetrates about 20 cm).

3. Test sequence

In the design of the test sequence, special attention needs to be given to Stage 2, i.e., the creation of a region around the well near residual CO₂ saturation. (Note that the inverse modeling approach applied here does not require that the field in Stage 3 is precisely at residual saturation, because the observed system behavior is sensitive to \( S_{gr} \) for all saturations greater than \( S_{gr} \). Nevertheless, we propose that the system state be close to residual saturation to maximize the sensitivity and to reduce the amount of extrapolation needed.) We have considered two approaches to creating a residual gas field: the first approach is to produce the mobile gas back and leave the trapped gas in the formation. This approach may be impractical or fail, because withdrawal of CO₂ slows down asymptotically as the residual gas saturation is approached, requiring the production of substantial brine volumes, into which the trapped CO₂ may gradually dissolve. Long testing time, the production of a large amount of brine, and the uncontrollable dissolution of CO₂ with an uncertain final saturation state make this approach unfeasible for our purposes. As an alternative, water saturated with CO₂ can be injected after CO₂ injection, pushing away the mobile gas from the well vicinity and leaving behind trapped CO₂ at its residual saturation. The co-injection of a small amount of CO₂ is essential to prevent CO₂ dissolution into the liquid phase. The proposed test sequence of each stage is summarized in Table 1 and Fig. 2.

The duration of each test as well as injection and withdrawal rates are site specific and should be chosen to maximize the information content of the data collected within operational constraints. The volume of reservoir interrogated by each test is also site specific and different for the three test types. In general, the smallest scale is investigated during the temperature test (1–2 m), a medium scale is investigated by the tracer test (tens of meters), and the largest scale is investigated by the pressure test (tens to hundreds of meters). This variation in spatial scale is a function of the diffusivity that governs each process. Given that the actual gas saturation distribution around the well is unknown, and that an effective, average parameter capturing residual gas trapping is sought, our approach of combining data from three test types that are based on different physical properties and predicated on different support scales adds to the overall robustness and representativeness of the estimated residual gas saturation parameter.

In this paper we explore through numerical simulations the time-series data of pressure, temperature, and tracer concentration that would be acquired during field studies. Based on synthetic data sets generated using a heterogeneous reservoir (modeled after con-

<table>
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<th>Table 1</th>
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<td><strong>Test sequence.</strong></td>
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<td><strong>Test</strong></td>
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<td><strong>Stage 1: reference tests</strong></td>
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<td>Thermal</td>
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<td>Hydraulic</td>
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<td>Pull-back</td>
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<tr>
<td>Tracer</td>
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<tr>
<td>Pull-back</td>
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<td><strong>Stage 2: creation of system with residual gas</strong></td>
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<tr>
<td>Optional thermal</td>
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<tr>
<td>CO₂ injection</td>
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<tr>
<td>Shut-in</td>
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<tr>
<td>Water injection (CO₂ saturated)</td>
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<tr>
<td>Shut-in</td>
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<tr>
<td><strong>Stage 3: characterization tests</strong></td>
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<tr>
<td>Thermal</td>
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<tr>
<td>Hydraulic</td>
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<td>Pull-back</td>
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<td>Tracer</td>
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<td>Pull-back</td>
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ditions at Otway), we will infer entrapped gas volume and residual gas saturation.

4. Analysis of proposed test design

This section includes two numerical analyses to examine the ability of the proposed approach to identify the residual trapping mechanism and its key parameters. The first analysis is a synthetic case, in which we consider various sources of uncertainty and demonstrate that the joint use of multiple data sets and joint use of reference and characterization tests improves the estimate of $S_{gr}$. For this analysis, a numerical model accounting for supercritical CO$_2$ injection into a heterogeneous brine formation is used to generate a set of synthetic observation data for sensitivity and inversion analyses. Gravity is explicitly accounted for. We will present the analysis results from simulations using a 3D model. Analyses are also conducted using a 2D model. The comparison of the two sets of results can help us to determine the appropriateness for using a 2D model for design calculations. In the second analysis, we discuss the approach in the context of the Otway CO$_2$ injection pilot test design.

Simulations are conducted using the EOS7C module of TOUGH2 (Pruess et al., 1999), a numerical simulator for non-isothermal multiphase, multicomponent flows in porous media. EOS7C (Oldenburg et al., 2004) was developed to model gas mixtures of methane and a non-condensible gas (in our case: CO$_2$), with one tracer present. This EOS module is incorporated in the iTOUGH2 (Finsterle, 2004) simulation–optimization code in preparation for formal sensitivity analyses and parameter identification. We use the inverse modeling capabilities of iTOUGH2 to calculate sensitivity and covariance matrices. The focus of this study is to demonstrate how multiple data sets and multiple tests can be jointly used to estimate $S_{gr}$ and its uncertainty.

Hysteresis is not included in the current study, even though it is expected that hysteresis in the capillary pressure and specifically the relative permeability functions plays a key role in the plume behavior and CO$_2$ trapping for an actual injection scenario (Doughty, 2007), and also during the proposed single-well test with its alternating drainage and imbibition sequence. However, while the absolute amount of CO$_2$ trapped depend on the subtleties in the hysteretic model, the design calculations discussed here are mainly concerned with relative changes in the variables observed during the well test. We acknowledge that accounting for hysteresis will lead to more complex process simulations and additional parameters that need to be determined, which in turn lead to higher estimation uncertainty. Examination of this more realistic, hysteretic system is subject of future research and may need to be considered when the proposed approach is applied for the analysis of actual field data.

4.1. Synthetic case analysis

The main purposes of this analysis are to investigate how various uncertainties affect the estimate of $S_{gr}$ and how the
joint use of different types of data and tests can help constrain the final estimate. For these purposes, we present analysis results conducted using a 3D model. CO₂ is assumed to be injected into a 5 m thick reservoir. The lateral extent of the model is 1 km. Geostatistical simulation techniques implemented in iTOUGH2 (Finsterle and Kowalsky, 2008) are used to generate a heterogeneous field of log-permeability modifiers with a correlation length of 20 m in the horizontal direction and 1 m in the vertical direction. The capillary-strength parameter \( \alpha \) of the van Genuchten constitutive relationships (Genuchten, 1980) is linked to the heterogeneous permeability field \( k \) according to the Leverett scaling rule, \( \alpha_2 = \alpha_1 \sqrt{k_2/k_1} \) (Leverett, 1941). A set of synthetic observation data for the test sequence described above is generated using the parameter values listed in Table 2.

Fig. 3 shows liquid saturation distributions after residual gas field is created (i.e., at 24 days) at three different elevations and for two vertical cross sections. We observe that CO₂ distribution is highly non-uniform, with more CO₂ at the top layer due to buoyancy effects, and horizontal patterns that differ significantly from radial symmetry of a homogeneous system. CO₂ saturation may be locally higher even relatively far away from the injection point as a result of heterogeneity. For example, the lowest liquid saturation seen in the top layer occurs away from the injection interval at a location with a high permeability, where gas enters relatively easily due to the corresponding weak capillary pressure (a correlation accounted for by the Leverett scaling).

In addition, we conduct similar analyses using a 2D radial model. Comparison of the 2D and 3D results help identify the relative importance of the heterogeneity structure (a realistic 3D field vs. a simplified, radially symmetric property distribution) on the evaluation of the test design.

4.1.1. Uncertainty quantification

Two types of uncertainty are considered: parametric uncertainty and uncertainty from geological heterogeneity. In addition to \( S_{gr} \), 11 parameters are considered uncertain, including hydrological parameters, thermal parameters, transport parameters, geological parameters and operational parameters. Table 2 contains the parameter values for which the sensitivity and uncertainty analyses are performed, along with their expected variation, which is used to normalize sensitivity coefficients. At the design stage, it is important to consider the major sources of uncertainty, maintaining the model complexity at a level appropriate for the intended analysis. Therefore, although residual liquid saturation \( (S_{lr}) \) is correlated to rock permeability and porosity (Holtz, 2006), residual liquid saturation effects are represented by a single, albeit uncertain parameter. Geological uncertainty refers to the random component of the spatial variability, i.e., the unknowable details of the heterogeneous permeability field. Note that irreducible geological uncertainty exists even though the geostatistical parameters of the permeability field may be known. We will first only consider parametric uncertainty, then add in geological uncertainty to examine its effect on the estimation uncertainty of \( S_{gr} \).

4.1.1.1. Parametric uncertainty. To investigate parametric uncertainty, sensitivity analyses with different combinations of data presumed available are performed. A dimensionless sensitivity coefficient is calculated as a relative measure of information content of a single observation point. It is defined as the ratio of change of the model output (i.e., the observable system response at the point in space and time where data are to be collected) over the change of the model input (i.e., the uncertain parameter to be estimated), scaled by the assumed measurement noise and expected parameter variation. Results of each type of output data at the two stages to the 12 uncertain parameters are plotted in Fig. 4. The sum
of the absolute values of the sensitivity coefficients for each data set indicates how much information each test contains with respect to the parameters.

While high sensitivity is necessary to determine the parameters of interest, strong correlations among the parameters may nevertheless lead to a large variance in our parameter estimates. A high estimation uncertainty (defined as the variance of the estimated parameter) is revealed by the diagonal elements of the parameter covariance matrix. Recalculating the parameter covariance matrix, which is calculated based on a linearity and normality assumption. Recalculating the parameter covariance matrix using different synthetic data sets indicates how estimation uncertainty may be reduced by using multiple data types and multiple tests. The estimation uncertainty of \( S_{gr} \) is likely to be small if we assume that all the other parameters are perfectly known. To avoid reaching an overly optimistic conclusion, it is essential to include all relevant parameters in the analysis to account for correlations of these parameters to \( S_{gr} \), which may significantly increase its estimation uncertainty to a more realistic value. To demonstrate how each kind of observation helps constrain the estimate of \( S_{gr} \), the estimation covariance matrix is calculated assuming that (1) only data from Stages 2 and 3 are available (i.e., no reference tests were performed); (2) only one type of data is available (i.e., either pressure, temperature or tracer data); (3) two types of data are available (any combination of pressure, temperature or tracer data); and (4) all three data types are available. Data listed in (2)—(4) include data from all three stages of the test sequence. The standard deviations of each uncertain parameter are listed in Table 3.

Although some of the geostatistical parameters are considered uncertain, the random component of the spatial variability referred to as geological uncertainty was removed by employing the same seed number for the creation of the permeability field in the linear error analysis as that used to generate the synthetic observation data. (Geological uncertainty is analyzed in the next subsection.) Information from Fig. 4 and Table 3 are complementary to each other. Fig. 4 shows that pressure data P4 in the characterization tests have a large sensitivity to the parameter of interest, \( S_{gr} \). However, pressure also has the largest sensitivity to almost all the other uncertain formation parameters with a few exceptions: dry heat conductivity is most sensitive to T2 (temperature data at maximum CO2 saturation); wet heat conductivity is most sensitive to T1 (temperature data under single-phase liquid conditions); Henry’s coefficient is most sensitive to C2 (tracer data in the characterization test), etc. Therefore, if a different pressure signal were obtained during the characterization tests, it is very difficult to determine which parameter is responsible for it. As a result, if only pressure data from the characterization tests were used, the estimation uncertainty for \( S_{gr} \) would be 0.08. While estimation uncertainty would be at a similar level (0.06) if only temperature data were used, the uncertainty would be much higher if only concentration data were used. This is partly because fewer concentration calibration data are available as none are collected in Stage 2 (unlike pressure and temperature data). However, the sensitivity of tracer data could be improved by using two dissimilar tracers, which result in more calibration points. When two types of data are used jointly, the estimation uncertainty is reduced due to the reduced parameter correlation; e.g., if both pressure and temperature data are available, the estimation uncertainty is reduced to 0.02. If all three types of data are used, this number is further reduced to 0.017.

Data from the reference tests (P1, T1 and C1) contain valuable information about the system under single-phase conditions. Including these data in the inversion helps determine single-phase parameters and other formation characteristics (e.g., heterogeneity). This independent estimation thus reduces the correlations among the parameters once two-phase conditions are invoked by gas injection, or lessens the negative impact of these correlated parameters on the estimate of \( S_{gr} \). This can be clearly demonstrated for permeability (log10 (k)), whose estimation uncertainty is reduced by 50% if reference tests are performed.

### Table 3

Estimation uncertainties assuming different available data sets.

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Without reference tests</th>
<th>With reference tests</th>
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<tr>
<td></td>
<td>All</td>
<td>P+T</td>
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<tr>
<td>( S_{gr} )</td>
<td>0.03</td>
<td>0.08</td>
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<tr>
<td>( S_c )</td>
<td>0.08</td>
<td>0.20</td>
</tr>
<tr>
<td>( \log_{10} (k) )</td>
<td>0.06</td>
<td>0.30</td>
</tr>
<tr>
<td>( \phi )</td>
<td>0.02</td>
<td>0.15</td>
</tr>
<tr>
<td>( n ) parameter</td>
<td>0.39</td>
<td>1.73</td>
</tr>
<tr>
<td>( \log_{10} (c_p) )</td>
<td>0.79</td>
<td>1.28</td>
</tr>
<tr>
<td>( \lambda_{dry} )</td>
<td>0.20</td>
<td>9.93</td>
</tr>
<tr>
<td>( \lambda_{wet} )</td>
<td>0.12</td>
<td>7.81</td>
</tr>
<tr>
<td>( \log_{10} (H^{-1}) )</td>
<td>0.09</td>
<td>32.3</td>
</tr>
<tr>
<td>( aa )</td>
<td>1.11</td>
<td>2.39</td>
</tr>
<tr>
<td>( \eta_i )</td>
<td>0.04</td>
<td>0.40</td>
</tr>
</tbody>
</table>

permeability uncertainty and other related parameters lead to a reduced estimation uncertainty of $S_{gr}$.

The values in Table 3 provide design guidelines rather than quantitative predictions of final estimation uncertainty. The following issues have to be considered: (1) the sensitivity and uncertainty analyses are performed at a single point in the parameter space. Given the nonlinear nature of the forward and inverse problems, different values will be obtained for other parameter combinations. (2) The estimation uncertainties listed in Table 3 may be too optimistic, because they are based on the assumption that the observed data can be matched to within measurement precision. Furthermore, conceptual modeling errors (e.g., incorrect relative permeability and capillary pressure functions, an incorrect geostatistical model) are not accounted for in this analysis. Nevertheless, these uncertainty analyses provide information on the relative information content of each data type and data set. Table 3 demonstrates that the joint use of multiple data types and the inclusion of reference test data will substantially increase our ability to determine the residual gas trapping capability of a CO2 storage formation.

4.1.1.2. Geological uncertainty. To study the impact of geological uncertainty on parameter estimation errors, multiple synthetic inversions are performed using 17 permeability fields created by varying the seed number. The hydrological, thermal, transport, geostatistical and operational parameters are estimated concurrently using both data from reference tests and characterization tests. Noise is added to observation data to mimic potential measurement errors. The geological uncertainty is calculated based on the sample statistics of the estimates from these inversions. Fig. 5 shows two examples of the permeability field that are used in the quantification of the geological uncertainty. The estimation uncertainties considering geological variability are listed in the last column of Table 3. When geological uncertainty is considered, the estimation uncertainties for all parameters are larger than those obtained above without the inclusion of geological uncertainty; specifically, the estimation uncertainty for $S_{gr}$ is about 0.03. It is important to point out that this estimation uncertainty is still optimistic since (1) the applied geostatistical model may not correctly represent the heterogeneity field; (2) there are additional uncertain parameters that are assumed to be perfectly known in the analysis. To reduce geological uncertainty, (1) vertical measurements from the well logs, or, if possible, measurements performed on recovered core should be used as conditioning points during the generation of the permeability field; and (2) geophysical measurements and imaging technologies could be employed during CO2 injection and integrated into the inversion analyses. This will reduce the geological uncertainty in the vicinity of the injection area.

While the uncertainty in the estimation of $S_{gr}$ has increased when accounting for geological uncertainty, the joint interpretation of the combined hydraulic, thermal, and tracer tests provides a reasonable design basis for which to assess the efficacy of residual trapping. Accounting for geological uncertainty in the estimation of $S_{gr}$ has thus been demonstrated to be essential to providing a realistic assessment of potential estimation errors, and can be accomplished by performing stochastic inversions.

4.1.2. 2D radial model vs. 3D model

The 3D analyses discussed above were repeated using a 2D radial model. Because of the forward process model needs to include non-isothermal, multiphase flow physics for correctly representing the dynamic behavior of the simulated system, it is much more difficult and time-consuming to perform an inverse analysis using a 3D model compared to a 2D model. However, a 2D radial model misrepresents heterogeneity (i.e., ring-shape heterogeneity). The goal of this subsection is to investigate the possibility of using a 2D radial model for a more efficient analysis by comparing its results to those of the 3D model.

In general, the estimation uncertainty results obtained from a 2D radial model are in qualitative agreement with the results from a 3D model without considering geological uncertainty. However, after geological uncertainty is added, the estimation uncertainty increases more for the 2D radial model. To summarize this comparison, we found that when considering pressure, temperature, and tracer tests combined with geologic uncertainty the residual gas saturation estimation uncertainty is 0.07 for the 2D radial model and 0.03 for the 3D model. This higher uncertainty for the 2D radial model is due to the fact that heterogeneity is encountered in serial mode in the dominant (radial) flow direction, increasing the impact of low-permeability regions. During the design stage, when the simulation purpose is to find the information content of potential observation data, it is sufficient to use a 2D radial model. However, if the goal is to estimate site-specific parameters, it is important...
to correctly represent heterogeneity of the system based on prior knowledge, thus a 3D model should be used.

4.2. Application to Otway Project

Recognizing the importance of saline aquifers for CO2 storage, the CO2CRC Otway Project proposed to conduct a single-well test aimed at better understanding residual trapping mechanisms and the related formation parameters. A pre-determined volume of CO2-rich gas (80% CO2 and 20% CH4 by volume; or 90% CO2 and 10% CH4 by mass) will be injected at a supercritical state into the Paaratte Formation. The composition of the injection gas stream is the result of the availability of naturally sourced gases from the nearby Buttress-1 well. The proposed injection formation is a sandy horizon in the Parratte Formation, at a depth of 1435 m (subsurface). There are two lithofacies in this zone: sand (59%) and shale (41%). For the numerical design calculations, the formation is considered heterogeneous and anisotropic, with porosity and permeability correlated to the facies type and conditioned on property values measured down-hole. The average horizontal and vertical sand permeabilities are 1 darcy and 0.01 darcy, respectively; the average porosity is 0.28. All the shale layers are considered to have 1 milli-darcy permeability.

4.2.1. Model description

The injection zone was modeled as a 2D radial system, extending 1000 m in the lateral direction and 26 m in depth. A heterogeneous permeability field is created using an indicator-based algorithm from the geostatistical software library GSLIB in iTOUGH2. The permeability field is generated using a variogram based on data measured at a nearby borehole, and conditioned on the lithologic data and permeability data from the well logs. The correlation length used in the variogram for sand and shale is 400 m and 300 m, with an anisotropy ratio of 0.0125 and 0.006, respectively. In the well data analysis, the bin size used for “bed thickness” of each lithofacies is 1 m; therefore, the maximum vertical discretization in the model is 1 m. The residual liquid saturation used in the simulation is 0.16. The well is perforated between −7 and −12 m. Vertical discretization of the model is 0.5 m for the upper 14 m and 1 m for the lower 12 m.

In the simulations the injected gas mixture is similar to the CO2 and CH4 ratio in the naturally sourced gas at Otway from the Buttress-1 well and ignores the higher hydrocarbons which are a small fraction of the gas stream. Supercritical CO2 is injected at a rate of 1.57 kg s\(^{-1}\) for 2 days, with CH4 co-injected at a rate of 0.16 kg s\(^{-1}\). Water injection rates during hydraulic and tracer testing as well as during the Stage 2 CO2 displacement phase is 1.74 kg s\(^{-1}\); fluid is produced at a rate of 0.4 kg s\(^{-1}\). During thermal testing, the borehole is heated with an intensity of 40 W m\(^{-1}\). Krypton (\(^{84}\)Kr) and xenon (\(^{132}\)Xe) are used as inert tracers that partition between the liquid and CO2-rich phases. These parameters were chosen as being operationally feasible using equipment available for the planned test.

4.2.2. Results

The proposed test sequence (see Fig. 2 and Table 1) was simulated assuming residual gas saturations of \(S_r = 0.1\), \(S_r = 0.2\), and \(S_r = 0.35\). Pressure, temperature, and tracer concentrations at the wellbore are recorded and plotted in Figs. 6–8. While the differences between these three cases are pronounced, one has to keep in mind that only a single hydraulic, thermal, and tracer response will be available from the actual tests, i.e., the challenge is to infer the unknown residual gas saturation \(S_r\) from a single set of observed pressure, temperature, and tracer concentration curves.

We first look at the system behavior throughout the model domain before we discuss each type of observation data. Fig. 9 shows the gas saturation for a system with \(S_r = 0.2\) at 15 days (immediately after CO2 injection) and 25 days (immediately after water injection). At day 15, the system reaches the highest CO2 saturation. Some of the injected gas moves upward due to buoyancy effects. At day 25, CO2 has moved further away, and the CO2 plume approached the residual saturation of 20%.

Fig. 6. Pressure responses at the wellbore as a function of time for different residual gas saturations.

Fig. 7. Temperature responses at the wellbore as a function of time for different residual gas saturations.

Fig. 8. Tracer breakthrough curves for different residual gas saturations.
4.2.2.1. Pressure responses. Fig. 6 shows the wellbore pressure as a function of time as calculated for the three residual saturation values. Pressures increase four times during the test sequence in response to (1) water injection into the fully water-saturated system (II–I4); (2) CO2 injection into the fully water-saturated system (drainage); (3) water injection to displace the CO2 (imbibition), leaving behind a CO2 plume near residual saturation; and (4) water injection into the residual CO2 plume (imbibition). Pressures decrease twice (D1 and D2), i.e., during (1) water production in Stage 1; and (2) production of water and CO2 during Stage 3 hydraulic and tracer testing. The pressure responses shown in Fig. 6 can be interpreted as follows:

(1) A comparison of the data from P1 and P2 is useful to infer the residual gas saturation. As discussed above, the curves calculated with different residual gas saturations show the sensitivity of the pressures response to the parameter of interest. Since only one curve will be available for data analysis, information about Sgr must be extracted from different time periods of the same curve, as the high sensitivity to residual saturation cannot be directly exploited.

(2) Pressure is only marginally sensitive to residual gas saturation during the CO2 injection phase. The reason for this relative lack of sensitivity is that during gas injection into the water-saturated formation, the total mobility is determined by that of the displaced water phase at full liquid saturation, and the invading gas phase at high gas saturation, both of which being relatively insensitive to Sgr. If hysteresis were accounted for, the sensitivity would disappear entirely, as the capillary pressure and relative permeabilities follow the primary drainage curves, which do not depend on the residual gas saturation.

(3) The pressure difference between the first and second pressure increases (I1 and I2) is due to the mobility and density contrasts between water and supercritical CO2.

(4) The third and forth pressure increases (I3 and I4) are almost the same because the pressure buildups of these water injection tests are determined by the mobility of the wetting phase near residual gas saturation, which is the same in both cases.

(5) The difference in pressures between the first production (from a fully water-saturated system) and second production (from the formation with trapped CO2) can also be used to infer Sgr, because the pressure decrease in the second test is determined by liquid mobility at Sgr.

4.2.2.2. Temperature responses. Fig. 7 shows the temperature responses at the wellbore. The three temperature increases correspond to the three heating periods. The last thermal test, conducted with trapped CO2 near the wellbore, shows some sensitivity to the residual gas saturation. The first temperature increase contains information on the wet heat conductivity; the second (when CO2 is at its maximum saturation) contains information on the dry heat conductivity; together, they can be combined with the third temperature increase to infer Sgr.

4.2.2.3. Tracer breakthrough curves. Tracer breakthrough curves are plotted in Fig. 8. Two noble gases are used as tracers in our simulations: krypton (84Kr) and xenon (133Xe). The Henry's coefficients are $3.86 \times 10^9$ Pa for Kr and $2.63 \times 10^9$ Pa for Xe at $65^\circ$C. By combining the information from two comparisons: (1) comparison of the BTC from the fully water saturated system and the BTC from the residual gas field; and (2) arrival time and tail differences between two different tracer BTCs from the residual gas field, the residual gas saturation can be inferred.

If only one tracer were used in the test, a highly volatile tracer (such as SF6) might not be suitable, because a large fraction of a highly volatile tracer would partition into the trapped gas phase, drastically reducing the amount of tracer recoverable in the liquid phase. If multiple partitioning tracers were used, large differences in Henry's coefficient (e.g., the highly volatile tracer SF6 and the highly water soluble refrigerant r134a (1,1,1,2 tetrafluoroethane)) will improve sensitivity to residual CO2 saturation. The use of any particular tracer needs to be supported by laboratory determination of a partitioning coefficient between the gas and water phase at the appropriate reservoir pressure and temperature.

5. Conclusion

In this paper, we present a design approach using sensitivity analyses and synthetic inversions to estimate effective residual CO2 saturation from single-well tests. Developing a robust single-well testing methodology is attractive because it obviates the need for drilling multiple wells at the great depths required for geological CO2 storage. Our design approach aims at addressing the challenges for estimating Sgr using single-well tests. We propose to (1) conduct a process-based data analysis for accurately representing the multiphase-flow process of the design sequence and determining the information content of observational data and the estimation uncertainty of the parameters of interest; (2) jointly invert Sgr using multiple types for reducing the correlations among various parameters; and (3) conduct reference tests under single-phase conditions for reducing the potential bias caused by formation heterogeneity.

To perform the design calculations, we built a physical-based model for a synthetic example and conducted a formal analysis for sensitivity and uncertainty quantification. Both parametric uncertainty and geological uncertainty were considered in the analysis. In addition to Sgr, 11 parameters including hydrological parameters, thermal parameters, transport parameters, geostatistical parameters, and operational parameters were considered to be uncertain. Results from the modeling analysis indicate that joint data analy-
sis provides sufficient sensitivity and complementary information to identify $S_{gr}$. Although pressure data from the characterization test are most sensitive to $S_{gr}$, as well as many other uncertain parameters, we are not able to obtain a unique estimate of $S_{gr}$ using pressure data alone due to correlations among parameters. Combining temperature data and tracer concentration data with pressure data reduces the estimation uncertainty. In addition, the differences between respective tests under single-phase and two-phase conditions reduce ambiguity and correlations among uncertain parameters. Finally, geological uncertainty affects data interpretation, and thus needs to be taken into account when estimating $S_{gr}$ and its uncertainty. Given inherent uncertainty caused by subsurface heterogeneity and measurement errors, the joint inversion of all three data sets and stochastic inversion are proposed as a practical approach to best constrain $S_{gr}$.

Our analyses are performed using both a 2D radial model and a 3D model. The 2D radial model has the advantage of being computationally less intensive; however, the radial symmetry leads to an unrealistic heterogeneity pattern. Nevertheless, the two models are in qualitative agreement if they are used to study information content and parametric uncertainty, even though geological uncertainty seems to be more pronounced when a 2D model is used. This indicates the potential appropriateness of using a 2D model at the design stage keeping in mind these limitations.

The synthetic studies provided insights that helped to design a practical field test to be carried out as part of the CO2CRC Otway Project. While the multiple phases of the test are operationally complex, the ability to use a single-well test for accurate appraisal of residual trapping mechanisms can provide considerable overall savings over methods that require multiple boreholes to meet the same objective.

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