Title
THE EQUILIBRIUM QUOTIENT FOR THE ISOMERIZATION OF BISULFITE ION FROM HSO3- TO SO3H-

Permalink
https://escholarship.org/uc/item/7vk7p7b8

Authors
Homer, D.A.
Connick, R.E.

Publication Date
1985-08-01
THE EQUILIBRIUM QUOTIENT FOR THE ISOMERIZATION OF BISULFITE ION FROM HSO$_3^-$ TO SO$_3^-$

D.A. Horner and R.E. Connick

August 1985
DISCLAIMER

This document was prepared as an account of work sponsored by the United States Government. While this document is believed to contain correct information, neither the United States Government nor any agency thereof, nor the Regents of the University of California, nor any of their employees, makes any warranty, express or implied, or assumes any legal responsibility for the accuracy, completeness, or usefulness of any information, apparatus, product, or process disclosed, or represents that its use would not infringe privately owned rights. Reference herein to any specific commercial product, process, or service by its trade name, trademark, manufacturer, or otherwise, does not necessarily constitute or imply its endorsement, recommendation, or favoring by the United States Government or any agency thereof, or the Regents of the University of California. The views and opinions of authors expressed herein do not necessarily state or reflect those of the United States Government or any agency thereof or the Regents of the University of California.
The Equilibrium Quotient for the Isomerization of Bisulfite Ion
from $\text{HSO}_3^-$ to $\text{SO}_3^-$

David A. Horner* and Robert E. Connick*

Department of Chemistry and the Materials and Molecular Research Division,
Lawrence Berkeley Laboratory, University of California
Berkeley, California 94720

ABSTRACT

Separate peaks in the oxygen-17 NMR spectra of sodium bisulfite solutions provide direct evidence for the existence in solution of two isomers of bisulfite ion: one with the proton bonded to the sulfur atom ($\text{HSO}_3^-$) and the other with the proton bonded to an oxygen atom ($\text{SO}_3^-$). The equilibrium quotient for the isomerization reaction was determined from measurements of peak areas, and was found to be $4.9 \pm 0.1$ at 298 K in solutions of ionic strength 1.0 m. The more abundant isomer exchanges oxygen atoms with water more rapidly than does the other, and on this basis is identified tentatively as $\text{SO}_3^-$.

This work was supported by the Director, Office of Energy Research, Office of Basic Energy Sciences, Chemical Science Division of the U.S. Department of Energy under Contract No. DE-AC03-76SF00098.

*Current address: North Central College, Naperville, Illinois 60566
Two plausible structures for the bisulfite ion are

\[ \text{H-S-}^0_{\text{O}} \quad \text{and} \quad \text{S-}^0_{\text{O-H}} \]

\((\text{HSO}_3^-)\) \quad \quad \quad \quad \quad \quad \quad \text{(SO}_3^\text{H}^-)\).

The appearance of the H-S stretch in Raman spectra shows that the \(\text{HSO}_3^-\) isomer exists in both aqueous solutions and solid salts,\(^{1-3}\) while positive but less conclusive evidence from Raman spectra\(^4,5\) supports the existence of \(\text{SO}_3^\text{H}^-\) in aqueous solution. Ab initio molecular orbital calculations\(^6,7\) predict that the gaseous ions of the two isomers have comparable stability.

We have found that under certain experimental conditions the two isomers of bisulfite ion produce separate peaks in the oxygen-17 NMR spectra of bisulfite solutions, providing direct evidence of the existence of both isomers. The equilibrium quotient for the isomerization reaction was determined through measurements of the areas of the spectral peaks.

Experimental Section

Analytical reagent grade sodium metabisulfite, sodium chloride, hydrochloric acid, and sodium hydroxide from the Mallinckrodt Chemical Company were used as supplied. Aqueous sodium bisulfite solutions were prepared in a nitrogen atmosphere glove bag by dissolving weighed amounts of sodium metabisulfite and sodium chloride in 3.00 ml of deoxygenated water which was enriched to about 50 mole percent \(^{17}O\) and 40 percent \(^{18}O\), but having natural abundance hydrogen and deuterium. The water was obtained from Monsanto Research Corporation's Mound Facility, and was purified by distillation.
Enough sodium chloride was added to each solution to give a calculated ionic strength of 1.0 m at 298 K. The pH of each solution was measured with an Orion Research model 601A digital pH meter and a Beckman Altex model 531164 combination pH electrode, which was calibrated with standard buffers. Adjustments of the pH of each bisulfite solution were made by addition of small amounts of 6M HCl or 6M NaOH. The solutions were placed in 10 mm NMR sample tubes which were then sealed with pressure caps and Parafilm.

Oxygen-17 nuclear magnetic resonance spectra were acquired on a pulsed NMR spectrometer consisting in part of a Cryomagnet Systems 4.7 Tesla superconducting magnet (corresponding to an 17O resonance frequency of 27.4 MHz), a homebuilt receiver and transmitter, a Nicolet Instrument Corporation 293A' programmable pulser, a Nicolet Instrument Corporation 1180 Data System, and a broadband, multinuclear, 10 mm probe. All experiments were run without a field-frequency lock on samples spinning at roughly 20 rps. The sample temperature was controlled by passing cooled or room temperature nitrogen gas through a heater and thence past the sample. A relay device connected to the heater and to a thermocouple located directly below the sample provided temperature regulation to ± 0.2 K. The actual temperature of the sample was determined by measuring the chemical shift difference between the proton NMR peaks of an ethylene glycol sample contained in a sealed capillary placed within the 10 mm sample tube. The published calibration curve has an estimated uncertainty of ± 0.9 K.

Each spectrum was obtained by Fourier transforming the sum of between 5000 and 10000 free decays produced with the standard one-pulse method or with a two-pulse sequence (θ x - τ - θ x - acquire; θ ≤ 45°) which suppressed the strong solvent signal. In the two-pulse sequence the frequency of the pulses is set equal to the Larmor frequency of the nuclei in the solvent site, and the time
τ is set equal to half the inverse of the difference in frequency between the solute and solvent peaks, so that at the end of time τ the solute and solvent magnetizations are 180° out of phase with each other. The second pulse returns the solvent magnetization to the z-axis of the rotating frame of reference, but tilts the solute magnetization further toward the xy-plane.

To eliminate rolling baselines caused by pulse breakthrough a delay of approximately 200 µsec was inserted between the end of the pulse and the beginning of data acquisition. In the one-pulse experiments the exact duration of this delay was set equal to the inverse of the frequency difference between the water and one of the bisulfite peaks \(^{10}\) to minimize or eliminate the need for first order phase correction.

All chemical shifts were measured relative to the water peak in the same spectrum; downfield shifts were assigned positive values.

Calculations of the concentrations of the various S(IV) species in the solutions were carried out using values of the dissociation equilibrium quotient of water and the two ionization quotients of sulfurous acid: \(pQ_w = 13.79\), \(pQ_{a1} = 1.37\), and \(pQ_{a2} = 6.34\), \(^{11}\) which are appropriate for conditions of 1.0 M, 1.0 m, and 1.0 m ionic strength, respectively, and 298 K. The value of the equilibrium quotient for the dimerization of bisulfite ion to form \(S_2O_5^{2-}\) was estimated to be 0.082 from the data of Connick, Tam, and von Deuster \(^5\) by first obtaining the equilibrium quotient at \([Na^+] = 5.0 M\) and a particular concentration of S(IV) by interpolation of the data in their Table 4, then extrapolating the result to an ionic strength of 1.0 M by assuming that the slope of their plot of \(\log Q_d vs. \mu^{1/2}\) is independent of the S(IV) concentration and that the ionic strengths of their \([Na^+] = 5.0 M\) solutions were 5.0 M. The same value was used for all calculations since it changed only trivially over the range of S(IV) concentrations used, and the very small
temperature dependence was ignored. The ionization equilibrium quotients listed above are given in units of molarity; they were assumed to have the same values in units of molality. Hydrogen ion concentrations were calculated from the pH meter reading, assumed to equal \(-\log_{10}(a_{H^+})\), and an activity coefficient of 0.754, which is the mean activity coefficient of 0.01 M HCl in 1 m NaCl. This activity coefficient was considered to be a reasonable substitute for the unknown activity coefficient in the bisulfite solutions. (The pH meter reading will be referred to by the term "pH".) Concentrations at temperatures other than 298 K were calculated after first correcting the values of \(Q_w\), \(Q_{a1}\), and \(Q_{a2}\) for changes in temperature using the enthalpies of reaction 13.55, -4, and -3 kcal/mole, respectively. (These reaction enthalpies are for solutions of ionic strength 0.5, 0, and 0, respectively; the values at unit ionic strength were not available.)

The traditional definition of molality was thought to be unsatisfactory for this work because it yields different concentrations of S(IV) for two solutions prepared in exactly the same manner save for the use of water having different mole fractions of the oxygen isotopes. It was decided to employ a unit of concentration which would yield equal S(IV) concentrations for solutions prepared using an equal number of moles of S(IV) and equal volumes of water, and would correspond to the traditional definition of molality if the water used were of normal isotopic composition. The unit chosen, moles of S(IV) per 55.5 moles of \(H_2O\), is called "molality" and is indicated by the symbol \(m\) throughout this paper.
Results and Discussion

In the pH range from three to five the $^{17}$O NMR spectra of 0.2 m sodium bisulfite solutions contain two or three peaks: a large water peak (greatly diminished in the two-pulse spectra) and one or two small peaks, presumably due to oxygen in S(IV) solute species, located between 170 and 200 ppm downfield from the water. Variations in acidity (as well as temperature and S(IV) concentration) affected the widths and chemical shifts of the peaks as well as the number of resonances observed (Fig. 1). Below pH 3.6 only one peak, located at about 175 ppm, appeared in addition to the water peak. As the pH was increased above 3.6 a second solute peak appeared at about 195 ppm, but at pH 5 only a single solute resonance, located at 193 ppm, was observed.

The areas of the S(IV) peaks were measured. Rough measurements of the relative areas of the 175 ppm and 195 ppm peaks were made by comparing simulated spectra of two overlapping Lorentzian lines to those two-pulse spectra in which two S(IV) peaks appeared. Although exact fits were not obtained due to the presence of chemical exchange and phase differences between the two peaks, it was possible to determine that the ratio of the area of the 195 ppm peak to that of the 175 ppm peak was between 5 and 10. Precise measurements of the areas of the peaks at 193 ppm and 175 ppm, at high and low pH, respectively, were obtained from the one-pulse spectra containing only one S(IV) resonance, using the water peak in the spectrum as an internal standard. The area of each peak was taken to be the product of its height and halfwidth. A Lorentzian curvefitting computer program was used to obtain both the height and halfwidth of the water peak, but did not give precise results for the linewidth and height of the small solute peak because the shape of that portion of the spectrum consisting of the superposition of the solute
peak and a wing of the water peak was very sensitive to slight phase errors. Therefore, the width of the S(IV) peak was obtained by fitting a Lorentzian curve to the corresponding peak in the two-pulse spectrum of the same solution at the same temperature. The height of the S(IV) peak in the one-pulse spectrum was then determined in the following manner: the wing of the water peak was approximated by a non-horizontal straight line in the region of the solute peak. Using $T_2$ from the two-pulse spectrum, the line segment of this line whose endpoints were at the frequencies $\omega_0 + 1/T_2$ and $\omega_0 - 1/T_2$ was determined. ($\omega_0$ is the location of the peak maximum, and $1/T_2$ is the half width at half height.) This line segment was then positioned so that its endpoints coincided with the spectral trace. The height of the peak was taken to be twice the vertical distance from the midpoint of the line segment to the trace of the peak.

The peak areas were corrected for the effect of signal decay which occurred during the finite delay time between the end of the pulse and the beginning of data acquisition by multiplying the area of each peak by $\exp(\tau/T_2)$, where $\tau$ is the delay time.

Table 1 shows the results of the determination of peak areas. The area of the peak at 193 ppm is so large that it can be accounted for only by assigning it to bisulfite ion, which was by far the major solute species in the solutions studied. According to the calculations of the concentrations of the various S(IV) species over 90% of the S(IV) was present in the form of bisulfite ion, while $SO_2^-$, $S_2O_5^{2-}$, and $SO_3^{2-}$ accounted for no more than 3%, 7%, and 3%, respectively. Given this information and the fact that the chemical shifts of $SO_2^-$ and $SO_3^{2-}$ are 520 ppm and 210 ppm, respectively, the peak at 175 ppm cannot be assigned to either $SO_2^-$ or $SO_3^{2-}$. Because its area is proportional to $[S(IV)]$ and not $[S(IV)]^2$, the 175 ppm peak also cannot be
assigned to $S_{2}O_{5}^{2-}$.

On the basis of the measured areas of the S(IV) peaks we conclude that the peak at 175 ppm must be assigned to one of the isomers of bisulfite ion, and the peak at 195 ppm to the other. Oxygen exchange occurs between the two isomers at a rate which increases with increasing pH, resulting in the coalescence of the two bisulfite resonances at pH 5 to give a single resonance located at 193 ppm. This interpretation of the spectrum is consistent with the results of a study of the kinetics of the system, which also shows that the 195 ppm isomer exchanges oxygen with water much more rapidly than does the 175 ppm isomer. The rate of oxygen exchange between water and the 195 ppm isomer is relatively slow at pH 5 and increases with increasing acidity, so that below pH 3.5 the 195 ppm resonance is too broad to observe.

The assignment of each of the two bisulfite peaks to a particular isomer was made by comparing the rates of oxygen exchange between water and each of the two bisulfite species. Exchange rates were obtained by studying the broadening of the 175 ppm resonance and the water peak; it was found that the pseudo first order rate constant for oxygen exchange from the 195 ppm site to water is at least 500 times as large as the pseudo first order rate constant for the exchange from the 175 ppm site to water. The rate law for the exchange from the 195 ppm site to water was found to be first order in both hydrogen ion and bisulfite ion. The $SO_{3}H^{-}$ isomer offers the simplest mechanism for this process:

$$SO_{3}H^{-} + H^{+} = SO_{2} + H_{2}O,$$

in which the incoming hydrogen ion need only attack the oxygen on which the bisulfite hydrogen is located. The corresponding mechanism involving $HSO_{3}^{-}$ is
more complicated because it requires in addition the migration of a hydrogen from the sulfur atom to an oxygen atom, and therefore is expected to occur more slowly. On this basis the 195 ppm resonance is assigned to \( \text{SO}_3^\text{H}^- \).

The equilibrium quotient for the isomerization reaction

\[
\text{HSO}_3^- \rightarrow \text{SO}_3^\text{H}^- \\
Q_2 = \frac{[\text{SO}_3^\text{H}^-]}{[\text{HSO}_3^-]}
\]

was determined using the areas of the 175 ppm peak in the spectra of the pH 3 solutions and the coalesced peak at 193 ppm in the spectra of the pH 5 solutions. If the only S(IV) species present in the solutions were HSO\(_3^-\) and SO\(_3^\text{H}^-\), then the quantity \([\text{HSO}_3^-] + [\text{SO}_3^\text{H}^-])/[\text{HSO}_3^-]\) would be equal to the ratio of the area of the 193 ppm peak at pH 5 to the area of the 175 ppm peak at pH 3. However, between 3% and 10% of the S(IV) in each of the solutions was present in the form of SO\(_3^{2-}\), S\(_2\)O\(_5^{2-}\), and SO\(_2\), the 170 resonances of which could perhaps be coalesced with the resonance of one of the isomers of bisulfite ion, creating the possibility that the areas of the 175 ppm and 193 ppm peaks were not proportional to the HSO\(_3^-\) and total bisulfite ion concentrations, respectively. It was therefore necessary to make some assumption concerning which S(IV) species contribute to the areas of the two peaks. It was assumed that the area of the peak at 175 ppm was due only to HSO\(_3^-\), while the area of the peak at 193 ppm arises from all the S(IV) species, i.e. HSO\(_3^-\), SO\(_3^\text{H}^-\), SO\(_3^{2-}\), SO\(_2\), and S\(_2\)O\(_5^{2-}\). The SO\(_2\) resonance (at 520 ppm) and the SO\(_3^\text{H}^-\) resonances were coalesced over the entire pH range from 3 to 5 due to the rapid oxygen exchange which occurs through reaction (1). An investigation of the 170 NMR spectra of two very concentrated bisulfite solutions ([S(IV)] > 4 m, pH = 3.2 and 4.5), in which the S\(_2\)O\(_5^{2-}\) concentration is large enough to make a significant contribution to peak areas, revealed
that at high concentrations the $S_2O_5^{2-}$ and bisulfite ion resonances are coalesced. It was assumed that the $S_2O_5^{2-}$ and the $SO_3H^-$ resonances were also coalesced at lower $S(IV)$ concentrations.

The assumption that $SO_3^{2-}$ contributes to the area of the 193 ppm peak at pH 5 is justified by the coalescence of the 175 ppm and 195 ppm resonances at that acidity. The increase with pH of the oxygen exchange rate between $HSO_3^-$ and $SO_3H^-$, which produces the coalescence of the 175 ppm and 195 ppm resonances at pH 5, is a consequence of the increase with pH of the rate of the following step in the exchange mechanism:

$$ (HSO_3^-)^* + SO_3^{2-} \rightarrow (SO_3^{2-})^* + SHO_3^- $$

(3)

where $SHO_3^-$ refers to both isomeric forms of bisulfite ion. The $(SO_3^{2-})^*$ can then be converted to $(SO_3H^-)^*$ by addition of a hydrogen ion from $SO_3H^-$, $HSO_3^-$, or $H^+$. The necessary condition for coalescence of two resonances is that the sum of the pseudo first order rate constants for oxygen exchange between the two sites be much larger than the difference between the precessional frequencies in the two sites. From the resonance frequency differences and the relative rate constants obtained from the concentration ratios it is readily shown that the $SO_3^{2-}$ resonance will be coalesced with the two bisulfite resonances if the latter are coalesced at pH 5 due to rapid oxygen exchange occurring via reaction (3). Thus $SO_3^{2-}$ will contribute to the 193 ppm resonance. The assignment of the $SO_3^{2-}$ oxygen population to either the 175 ppm or the 195 ppm site is unimportant at low pH because the concentration of $SO_3^{2-}$ is very small at high acidity.

Under the aforementioned assumptions the equilibrium quotient for reaction (2) could be evaluated by converting the area of the 175 ppm peak to
a \( \text{HSO}_3^- \) concentration and comparing this concentration to the calculated sum of the concentrations of the two isomers for the same solution. It was thought best, however, to use the area of the 193 ppm peak as a measure of the total concentration of bisulfite ion, in the hope that any systematic error inherent in the measurement of peak areas would be at least partially cancelled by the use of area measurements to obtain both the \( \text{HSO}_3^- \) concentration and the total bisulfite ion concentration. The equilibrium quotient for reaction 2 was evaluated using the formula

\[
Q_2 = \frac{\text{peak area corresponding to total } \text{HSO}_3^- \text{ concentration at pH 3}}{\text{area of 175 ppm peak at pH 3}} - 1. \tag{4}
\]

The area corresponding to the total bisulfite ion concentration at pH 3 was obtained using the area of the 193 ppm peak at pH 5 in the following manner. The concentrations of \( \text{HSO}_3^- \), \( \text{SO}_2^- \), \( \text{SO}_3^{2-} \), and \( \text{S}_2\text{O}_5^{2-} \) were calculated for the pH 5 solutions, and the fraction of the 193 ppm peak area attributable to bisulfite ion was determined. The area corresponding to the total bisulfite ion concentration in the pH 3 solution was obtained by multiplying the area corresponding to the total bisulfite ion concentration in the pH 5 solution by the ratio of the calculated bisulfite ion concentration at pH 3 to that calculated at pH 5. To minimize error arising from the estimated value of the equilibrium quotient for the dimerization of bisulfite ion (which was used in calculating the bisulfite ion concentrations), a pH 5 solution having nearly the same bisulfite concentration as the pH 3 solution was always used.

The values of the parameters used in the calculation of \( Q_2 \) are listed in Table 2 together with the resulting values of the equilibrium quotient. Figure 2 shows a plot of \( \log Q_2 \) vs. \( 1/T \). A non-weighted linear least squares treatment of the data yields
\[ \ln Q_2 = (-3.232 \pm 0.5274) + (1438 \pm 151.2)/T \]  

(5)

as the equation of the best straight line through the points, with a covariance of -79.71 between the slope and y-intercept. From the values of the slope and intercept one obtains \( \Delta H_2 = -2.9 \pm 0.3 \) kcal/mole and \( \Delta S_2 = -6 \pm 1 \) cal K\(^{-1}\) mol\(^{-1}\).

The oxygen-17 NMR spectra of bisulfite solutions provide the most convincing evidence to date for the existence of the two isomers of bisulfite ion, HSO\(_3^-\) and SO\(_3^2-\), and allow the first measurement of their equilibrium concentration ratio.

Acknowledgements

This work was supported by the Director, Office of Energy Research, Office of Basic Energy Sciences, Chemical Science Division of the U.S. Department of Energy under Contract No. DE-AC03-76SF00098.
REFERENCES AND NOTES

(1) Simon, A. Angew. Chem. 1947, A59, 247.
(14) We observed an $^{17}$O chemical shift of 210 ppm for $SO_3^{2-}$ in 0.3 m $Na_2SO_3$.
(15) Connick, R. E.; Horner, D. A., to be published.
Table 1
Variation of S(IV) Peak Area with S(IV) Concentration and Acidity in $^{17}$O NMR Spectra of Sodium Bisulfite Solutions in Which Only One S(IV) Peak Appears$^a$

<table>
<thead>
<tr>
<th>[S(IV)]$^b$</th>
<th>pH meter$^c$ reading</th>
<th>S(IV) peak chemical shift, ppm</th>
<th>area S(IV) peak $\times \frac{55.5}{\text{area H}_2\text{O peak}}$</th>
<th>[S(IV)]$^d$</th>
</tr>
</thead>
<tbody>
<tr>
<td>0.201 m</td>
<td>4.98</td>
<td>193</td>
<td>3.24 ± 0.07</td>
<td></td>
</tr>
<tr>
<td>0.449 m</td>
<td>4.97</td>
<td>191-193</td>
<td>3.05 ± 0.04</td>
<td></td>
</tr>
<tr>
<td>0.101 m</td>
<td>3.01</td>
<td>175-178</td>
<td>0.50 ± 0.08</td>
<td></td>
</tr>
<tr>
<td>0.199 m</td>
<td>3.00</td>
<td>175-178</td>
<td>0.42 ± 0.06</td>
<td></td>
</tr>
<tr>
<td>0.456 m</td>
<td>3.02</td>
<td>173-177</td>
<td>0.42 ± 0.06</td>
<td></td>
</tr>
</tbody>
</table>

$^a$The ionic strength of each solution was 1.0 m.

$^b$The S(IV) was added as Na$_2$S$_2$O$_5$.

$^c$Measured at room temperature.

$^d$Averages of values obtained from spectra recorded between 273 K and 303 K, including a correction for the signal decay which occurred during the pre-acquisition delay time (see Experimental Section).
Table 2

Values of Parameters Used in the Calculation of $Q_2$, the Equilibrium Quotient for the Reaction $\text{HSO}_3^- = \text{SO}_3\text{H}^-$

<table>
<thead>
<tr>
<th>Ionic Strength 1.0 m</th>
<th></th>
</tr>
</thead>
<tbody>
<tr>
<td>[S(IV)]</td>
<td>pH meter reading</td>
</tr>
<tr>
<td>S(IV) peak</td>
<td>S(IV) peak</td>
</tr>
<tr>
<td>reading</td>
<td>chemical shift, ppm</td>
</tr>
<tr>
<td>[SHO$_3^-$ oxygen]</td>
<td>[total S(IV) oxygen]</td>
</tr>
<tr>
<td>T, °C</td>
<td>[SHO$_3^-$]</td>
</tr>
<tr>
<td></td>
<td>area S(IV) peak</td>
</tr>
<tr>
<td>0.201 m</td>
<td>4.98</td>
</tr>
<tr>
<td>0.199 m</td>
<td>3.00</td>
</tr>
<tr>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td></td>
</tr>
<tr>
<td>0.449 m</td>
<td>4.97</td>
</tr>
<tr>
<td>0.456 m</td>
<td>3.02</td>
</tr>
<tr>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td></td>
</tr>
</tbody>
</table>

$^a$Chemical shift, ppm

$^b$Oxygen content in ppm

$^c$Area S(IV) peak

$^d$Area H$_2$O peak

$^e$Molar error

$^f$Molar error
Table 2 (cont.)

a Measured at room temperature.

b Calculated using literature values of equilibrium quotients, as explained in the text.

Includes a correction for the effect of signal decay which occurred during the preacquisition delay time (see Experimental Section).

d See text for method used to calculate Q₂.

e Averages of values obtained from spectra recorded in the listed temperature range. The area of the 193 ppm peak should not have a significant temperature dependence.
Figure 1
Figure 2

\[
\frac{[\text{SO}_3\text{H}^-]}{[\text{HSO}_3^-]} \quad \text{versus} \quad 10^3/T \quad (K^{-1})
\]

- ○ \([\text{S(IV)}] = 0.20 \text{ m}\)
- ● \([\text{S(IV)}] = 0.45 \text{ m}\)
Captions for Figures

Figure 1. Traces of a portion of the $^{17}$O NMR spectrum of sodium bisulfite solutions of various acidities, showing how the number of peaks in the spectrum as well as the widths and chemical shifts of the peaks are dependent upon the pH of the solution. Every solution had a S(IV) concentration of 0.20 m and an ionic strength of 1.0 m. The temperature of each solution was between 277 K and 283 K. The room temperature pH meter readings are indicated. The $^{17}$O spectra were recorded at a radio frequency of 27.377 MHz using the two-pulse sequence described in the Experimental section. The chemical shift scale is one on which the water peak in the one-pulse spectrum of each solution has a chemical shift of zero. Downfield shifts are positive. The vertical expansion is not the same for different traces.

Figure 2. The temperature dependence of the equilibrium quotient $Q_2$ for the reaction $\text{HSO}_3^- = \text{SO}_3\text{H}^-$. The equilibrium quotient was determined at an ionic strength of 1.0 m for solutions having S(IV) concentrations of 0.20 m and 0.45 m. The straight line represents the non-weighted linear least squares fit of the data, and has the equation $\ln Q_2 = (-3.232 \pm 0.5274) + (1438 \pm 151.2)/T$, with a covariance of -79.71 between the slope and y-intercept.
This report was done with support from the Department of Energy. Any conclusions or opinions expressed in this report represent solely those of the author(s) and not necessarily those of The Regents of the University of California, the Lawrence Berkeley Laboratory or the Department of Energy.

Reference to a company or product name does not imply approval or recommendation of the product by the University of California or the U.S. Department of Energy to the exclusion of others that may be suitable.