Lawrence Berkeley National Laboratory

Title
Large Releases from CO2 Storage Reservoirs: A Discussion of Natural Analogs, FEPS, and Modeling Needs

Permalink
https://escholarship.org/uc/item/961519s7

Authors
Birkholzer, J.
Pruess, K.
Lewicki, J.L.
et al.

Publication Date
2008-05-22
Large Releases From CO₂ Storage Reservoirs:
A Discussion of Natural Analogs, FEPs, and Modeling Needs

Jens Birkholzer¹, Karsten Pruess¹, Jennifer Lewicki¹, Jonny Rutqvist¹, Chin-Fu Tsang¹, Anhar Karimjee²

¹Lawrence Berkeley National Laboratory, Berkeley, CA, USA
²U.S. Environmental Protection Agency, Washington D.C., USA

Abstract
While the purpose of geologic storage in deep saline formations is to trap greenhouse gases underground, the potential exists for CO₂ to escape from the target reservoir, migrate upward along permeable pathways, and discharge at the land surface. In this paper, we evaluate the potential for such CO₂ discharges based on the analysis of natural analogs, where large releases of gas have been observed. We are particularly interested in circumstances that could generate sudden, possibly self-enhancing release events. The probability for such events may be low, but the circumstances under which they occur and the potential consequences need to be evaluated in order to design appropriate site-selection and risk-management strategies. Numerical modeling of hypothetical test cases is suggested to determine critical conditions for large CO₂ releases, to evaluate whether such conditions may be possible at designated storage sites, and, if applicable, to evaluate the potential impacts of such events as well as design appropriate mitigation strategies.

Introduction
Large amounts of CO₂ would need to be injected into deep geological formations if geological sequestration is chosen as a viable option for reducing greenhouse gas emissions to the atmosphere. The extent of injected CO₂ plumes would be on the order of 100’s of km², making it likely that caprock imperfections are encountered allowing some CO₂ to escape from the intended storage formation. At typical temperature and pressure conditions, CO₂ is less dense than water, and therefore is buoyant in most subsurface environments. Thus, leaking CO₂ would have the tendency to migrate upwards into the shallow subsurface and ultimately to the atmosphere. Discharge at the land surface is not necessarily a concern, as CO₂ is a naturally abundant and relatively benign gas in low concentrations. However, there is a potential risk to health, safety, and environment (HSE) in the event that large localized fluxes of CO₂ were to occur. Large-magnitude releases of gas (e.g., CO₂, natural gas) from depth to the near-surface environment that have occurred in natural settings can serve as analogs for the potential release of CO₂ from geologic storage sites. Analysis of these analogs provides important insight into the features, events, and processes (FEPs) associated with the CO₂ release, as well as the needs for further evaluation and modeling analysis.

Methodology and Results
Various natural analogs with large releases of CO₂ have been evaluated, placing emphasis on the geologic model for CO₂ accumulation, processes leading to the releases of CO₂, pathways for migration, and type of release at the surface [1]. Table 1 gives a summary of these natural analogs and their key characteristics. Industrial analogs, which may include well blowouts during drilling of new wells or leakage along existing wells with insufficient plugs, have not been evaluated in this paper. While the large number of abandoned wells is a major concern for geological storage of CO₂ in depleted or near-depleted oil and gas reservoirs [2], they may be of smaller importance for deep saline formations, simply because deep wells are much less abundant. While detailed information on geologic models of natural CO₂ accumulation or migration pathways is often unavailable, a few general conclusions can be drawn from the analog summary in Table 1 with respect to the risk assessment of geologic storage of CO₂.

1. Unsealed faults and fracture zones are the main pathways for gas migration in most natural analogs with large CO₂ discharge. It is thus mandatory to be able to predict qualitatively and quantitatively the fast upward flow of CO₂ in high-permeability zones that may extend from depth to surface. During upward migration, CO₂ is affected by transitions from super- to sub-critical conditions, phase changes between liquid and gaseous CO₂, phase partitioning between water-rich and CO₂-rich phases, and adiabatic cooling as a result of pressure decline. A series of hypothetical, yet realistic scenarios with fast CO₂ pathways should be modeled, while attempting to capture as much “thermodynamics” as possible (see “CO₂ Migration Along a Fault”). Such
studies would serve to gain a better understanding of the fluid flow and heat transfer processes that would accompany CO$_2$ migration away from the primary storage reservoir, towards shallow depths and ultimately to the land surface. Some of these processes may be beneficial in that they prevent or retard upward migration (e.g., attenuation of CO$_2$ in multi-layer systems; energy losses through adiabatic cooling), while others may enhance CO$_2$ upflow (e.g., stress-induced increases in fault permeability; reduced phase interference).

2. Many natural releases of CO$_2$ have been correlated with a specific event that has triggered the release, such as seismic activity leading to geomechanical damage in sealing caprocks (e.g., hydraulic fracturing, fault slip and reactivation). The potential for processes to cause such damage and trigger the release of CO$_2$ from a storage reservoir should be evaluated with appropriate geomechanical modeling tools (see “Geomechanical Failure Analysis in a Multi-Layer System”).

3. CO$_2$ can both accumulate beneath and be released from primary (deep) and secondary (shallow) reservoirs with caprock units located at a wide range of depths. In general, a sequence of caprock units above the intended storage formation would be a desirable feature for a geological sequestration site, because it provides additional barriers for leaking CO$_2$. However, when considering the possibility of large, sudden discharges at the land surface, accumulation of CO$_2$ in a secondary formation in the shallow subsurface could actually have detrimental effects. Consider for example a shallow anticlinal structure with a low-permeability caprock, where CO$_2$ leaking from depth would slowly accumulate. If the accumulated CO$_2$ would then be released, as triggered, for example, by sudden geomechanical damage of the sealing caprock, the pathway to the land surface might be too short to allow for significant mitigation or retardation of the plume. CO$_2$ present as a separate gas phase would move upward by buoyancy forces and pressure differences (Figure 1a). CO$_2$ dissolved in water would be subject to rapid degassing following depressurization, which could result in fast-rising expanding bubbles (Figure 1b). The potential for shallow CO$_2$ accumulation should be considered in site evaluation efforts and, if applicable, the possible release of CO$_2$ from such secondary accumulation should be evaluated in numerical modeling studies.

![Figure 1](image.png)

Figure 1. Sudden CO$_2$ leakage from secondary accumulation in shallow reservoirs. (a) CO$_2$ accumulates as a separate phase. (b) CO$_2$ is dissolved in water. A fracture or fault zone may develop as a result of geomechanical damage or in response to seismic activity.

Simulation Examples

Here we briefly summarize results of numerical simulation studies as examples of the type of scenario modeling needed to further our understanding of CO$_2$ storage and its related risks. The first example is on CO$_2$ leakage along a continuous fault zone from depth to surface. The focus is here on capturing the complex thermodynamics in detail to see whether self-limiting and self-enhancing features would tend to slow or accelerate the upward migration of CO$_2$. The second example involves geomechanical modeling to determine the potential for fault reactivation and hydraulic fracturing in a multi-layered reservoir-caprock system.
CO$_2$ Migration along a Fault

Figure 2 shows a schematic model of a fault zone, along with simulation results for CO$_2$ discharge through this fault. The fault initially contains water in a normal geothermal gradient of 30 °C/km with a land surface temperature of 15 °C, in hydrostatic equilibrium. CO$_2$ discharge is initiated by injecting CO$_2$ at an overpressure of approximately 10 bar in a portion of the fault at 710 m depth. The numerical simulation includes two- and three-phase flow of an aqueous phase and liquid and gaseous CO$_2$ phases in the fault, as well as conductive heat transfer with the adjacent wall rocks that are assumed to be impermeable [5, 6].

![Figure 2. CO$_2$ leakage along a fault zone [5, 6]. A schematic model of a fault zone is shown on the left. The right panel gives temporal variation of CO$_2$ leakage fluxes at two different positions at the land surface. Total flow system volume with three-phase conditions is also shown.](image)

Strong cooling occurs due to the Joule-Thomson effect as rising CO$_2$ expands. Additional temperature decline is evident when liquid CO$_2$ boils into gas. The simulations show persistent flow cycling with increasing and decreasing leakage rates at the land surface after a period of initial growth. Monotonic behavior is observed when flow-system temperatures are held constant at their initial values. The cyclic behavior is explained in terms of varying fluid phase composition, due to heat transfer limitations, giving rise to an interplay between self-enhancing and self-limiting features. Overall, the cyclic leakage rates are consistently smaller than those obtained from the fixed-temperature simulation, showing that Joule-Thompson cooling and other thermodynamic effects generate an overall slowing of CO$_2$ migration (i.e., the upward flow of CO$_2$ is overall self-limiting).

Geomechanical Failure Analysis in a Multi-Layer System

In a hypothetical multi-layer system, CO$_2$ is injected for 30 years in a 200-meter thick permeable saline water formation located at 1600 meters depth (Figure 3). Several layers of caprock as well as water-bearing formations are located above the intended injection unit, all of which are intersected by a permeable fault zone. During injection, CO$_2$ migrates laterally and upwards in the intended storage formation, driven by injection pressure and buoyancy forces. When the plume encounters the fault zone intersecting the containing caprock unit, a considerable amount of CO$_2$ migrates upwards, spreads laterally into the upper overlying zones, and may cause considerable fluid pressure increase there. Based on the changes in effective stresses, the potential for fault slip and fracturing are calculated using failure criteria adopting conservative rock mass strength parameters [7].

![Figure 3. Model geometry for simulation of CO$_2$ injection into a multilayered reservoir-caprock system](image)
Figure 4 presents the potential for fault slip and hydraulic fracturing for two different initial stress regimes—a compressional stress regime (horizontal stress larger than vertical) and an extensional stress regime (horizontal stress smaller than vertical). Results are given in terms of pressure margins to the onset of shear slip or fracturing. A positive pressure margin in Figure 4 implies that the local fluid pressure may be above the critical pressure for onset of geomechanical damage. Dark contours indicate areas of the highest potential for onset of shear slip. Results suggest that, once leakage of CO$_2$ occurs, the potential for fault reactivation and fracturing could be larger in the overlying units than in the intended storage unit, a result of the smaller initial stresses in shallower units. Significant differences are observed between the compressional and extensional stress regimes. In the case of a compressional stress regime (Figure 4a), the shear slip is most likely to be initiated in subhorizontal fractures at the interface between the permeable formation layers and the overlying caprock unit. In the case of an extensional stress regime (Figure 4b), the shear slip is likely to occur in subvertical fractures in the uppermost aquifer and in the overburden. An extensional stress regime may also allow for hydraulic fracturing at the bottom of the uppermost caprock unit. Our analysis thus demonstrates that for evaluation of the maximum sustainable CO$_2$-injection pressure, it is essential to have a good estimate of the three-dimensional in situ stress field.

Figure 4. Calculated pressure margin for shear slip under (a) compressional stress regime and (b) extensional stress regime. The one and only location for hydraulic fracturing is also indicated in (b).

Conclusions

Natural analogs with significant CO$_2$ discharges at the land surface have been evaluated to identify FEPs and related modeling needs in support of risk assessment for geological storage of CO$_2$. We have discussed areas where further numerical analyses on hypothetical, yet realistic scenarios would be beneficial, and have presented example simulations.

Acknowledgment

We thank the U.S. Environmental Protection Agency, Office of Water and Office of Air and Radiation, for funding this study under an Interagency Agreement with the U.S. Department of Energy at the Lawrence Berkeley National Laboratory, contract number DE-AC02-05CH11231.

References

Table 1. Summary of natural large releases of CO₂.

<table>
<thead>
<tr>
<th>Site</th>
<th>CO₂ Source</th>
<th>Geologic model for accumulation</th>
<th>Event triggering leakage</th>
<th>Pathway for leakage</th>
<th>Type of release</th>
</tr>
</thead>
<tbody>
<tr>
<td>Mammoth Mountain, CA USA</td>
<td>Magmatic + thermal decomposition of carbonates</td>
<td>Accumulation at ~2 km depth in porous/fractured rock under caprock</td>
<td>Seismic activity and reservoir pressurization</td>
<td>Faults and fractures</td>
<td>Fast, diffuse, vent, spring</td>
</tr>
<tr>
<td>Solfatara, Italy</td>
<td>Magmatic + thermal decomposition of carbonates</td>
<td>Relatively shallow zone of fractured rock contains gas phase and overlies aquifers, then magma body at several km depth</td>
<td>No specific release event captured</td>
<td>Faults and fractures</td>
<td>Diffuse and vent</td>
</tr>
<tr>
<td>Mátrádereszkő, Hungary</td>
<td>Geothermal/copper-zinc mineralization</td>
<td>CO₂ accumulates in karst water reservoir (~1 km depth)</td>
<td>No specific release event captured</td>
<td>Faults and fractures</td>
<td>Diffuse, vent, spring</td>
</tr>
<tr>
<td>Láteralandera, Italy</td>
<td>Thermal decomposition of carbonates</td>
<td>CO₂ accumulates in liquid-dominated, carbonate geothermal reservoir capped by hydrothermally altered volcanics</td>
<td>No specific release event captured</td>
<td>Faults and fractures</td>
<td>Diffuse, vent, spring</td>
</tr>
<tr>
<td>Albani Hills, Italy</td>
<td>Magmatic + thermal decomposition of carbonates</td>
<td>Deep pressurized reservoirs in structural highs of sedimentary bedrock</td>
<td>Slow releases with several sudden large releases also occurring, possibly triggered by seismic activity</td>
<td>Faults and fractures</td>
<td>1995 and 1999 events Fast, diffuse, vent, spring/well</td>
</tr>
<tr>
<td>Dieng, Indonesia</td>
<td>Magmatic</td>
<td>Unknown</td>
<td>Volcanic, possibly “pneumatic”, eruptions</td>
<td>Fissure</td>
<td>Eruptive</td>
</tr>
<tr>
<td>Rabaul, Papua New Guinea</td>
<td>Magmatic</td>
<td>Unknown</td>
<td>Unknown</td>
<td>Fractures</td>
<td>Fast, vent</td>
</tr>
<tr>
<td>Lakes Monoun and Nyos, Camaroon</td>
<td>Magmatic</td>
<td>Accumulation in deep lake and stable stratification</td>
<td>Rapid lake turnover triggered at Monoun by landslide, Nyos trigger unknown</td>
<td>NA</td>
<td>Eruptive (limnic)</td>
</tr>
<tr>
<td>Laacher See, Germany</td>
<td>Magmatic</td>
<td>NA</td>
<td>Seasonal lake overturn and mixing</td>
<td>NA</td>
<td>Diffusive and bubbling from lake surface, diffuse from lake shore</td>
</tr>
<tr>
<td>Clear Lake, CA, USA</td>
<td>Thermal decomposition of metasedimentary rocks, minor magmatic component</td>
<td>CO₂ derived from liquid-dominated geothermal reservoir hosted in marine metasedimentary rocks</td>
<td>No specific release event captured</td>
<td>Faults and fractures</td>
<td>Gas vents, springs</td>
</tr>
<tr>
<td>Paradox Basin, UT, USA</td>
<td>Thermal decomposition of carbonates</td>
<td>Reservoirs are vertically stacked, sandstone units, in fault-bounded anticlinal folds, capped by shale/siltstone units</td>
<td>No specific release event captured</td>
<td>Faults and fractures</td>
<td>Diffuse, gas seeps, springs</td>
</tr>
<tr>
<td>Florina Basin, Greece</td>
<td>Thermal decomposition of carbonates</td>
<td>Reservoirs are vertically stacked, limestone and sandstone units (upper unit at 300 m depth), capped by silts and clays</td>
<td>No specific release event captured</td>
<td>Slow leakage along rock discontinuities</td>
<td>Springs, gas seeps</td>
</tr>
</tbody>
</table>