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Journal
Global Biogeochemical Cycles, 24(3)

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Publication Date
2010-07-31

DOI
10.1029/2009GB003670

Supplemental Material
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Effect of smoke on subcanopy shaded light, canopy temperature, and carbon dioxide uptake in an Amazon rainforest

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Received 11 September 2009; revised 4 March 2010; accepted 23 March 2010; published 31 July 2010.

[1] Daytime Net Ecosystem CO2 uptake (NEE) in an Amazon forest has been shown to increase significantly during smoky periods associated with biomass burning. We investigated whether the increase in CO2 uptake is caused by increased irradiance in the lower canopy, which results from increased above-canopy diffuse light, or by decreased canopy temperature, which results from decreased above-canopy net radiation. We used Sun photometers measuring aerosol optical depth to find nonsmoky (Aerosol Optical Depth (AOT) < 0.35), smoky (AOT > 0.5) and very smoky (AOT > 0.7) periods for the Tapajos region in the Amazon. Using a network of subcanopy photosynthetic photon flux density (PPFD) sensors, we detected a ∼4 μmol m−2 s−1 increase in subcanopy diffuse light during smoky periods relative to nonsmoky periods. Using a pyrgeometer to measure upwelling longwave radiation and, hence, canopy surface temperature, we found a ∼0.5°C cooling relative to air temperature during smoky periods. We modeled subcanopy irradiance based on the subcanopy PPFD sensors and combined this with subcanopy leaf photosynthesis measurements to determine how the increased lower canopy light affected NEE. We used the relationship between temperature and NEE measured by eddy covariance to determine the effect of decreased canopy temperature on canopy CO2 uptake. We found that the increase in CO2 uptake at high aerosol optical depths is primarily a result of increased shaded light in the subcanopy (accounting for ∼80%) and to a lesser extent the effect of decreased canopy temperature (accounting for ∼20%).


1. Introduction

[2] Extensive biomass burning in the Amazon rainforest creates atmospheric aerosols that can affect the region in multiple ways. Aerosols from tropical forest fires can reduce surface solar radiation by as much as 375 W m−2 and reduce surface temperatures by 2°C–3°C [Schafer et al., 2002]. Aerosols may cool the surface while warming the lower atmosphere (2–4 km) [Eck et al., 1998; Satheesh and Ramanathan, 2000]. Aerosols may indirectly warm the surface by reducing shallow cumulous clouds during periods of heavy smoke, thereby, causing an increase in solar radiation absorbed by the surface [Koren et al., 2004]. Aerosols can influence precipitation by changing the height at which clouds form [Andreae et al., 2004].

[3] Atmospheric aerosols from biomass burning may influence ecosystems directly by scattering light, which decreases overall irradiance but increases the diffuse component. Clouds and volcanic ash, in addition to biomass burning, can increase diffuse light, and their effect on ecosystems has been studied previously. Studies comparing Sun photometers measuring optical depth [Holben et al., 1998] to eddy covariance towers measuring CO2 exchange [Baldocchi et al., 2001] concluded that diffuse light increased CO2 uptake in crops and forests and decreased CO2 uptake in grasslands [Niyogi et al., 2004]. Other studies have shown an increase in canopy light use efficiency by forests during periods of increased diffuse light [Misson et al., 2005; Baldocchi, 1997; Knohl and Baldocchi, 2008]. Large volcanic eruptions increase diffuse light and provide a global test of the effect of diffuse light on gross production. An increase in aerosols from the Mount Pinatubo eruption coincided with a global increase in net CO2 uptake [Gu et al., 2003], possibly because of reduced respiration [Lucht et al., 2002]. However, in northern ecosystems, Net Primary Production (NPP)
decreased in response to volcanic events, based on an analysis of tree rings [Krakauer and Randerson, 2003].

[4] Most of these studies attribute the positive correlation between CO$_2$ uptake and aerosol optical depth to increased subcanopy diffuse light [Roderick et al., 2001]. Canopy top leaves are generally light saturated during midday conditions, and a marginal reduction in total irradiance does not decrease photosynthesis by these leaves. Subcanopy leaves, on the other hand, are often strongly light limited, and an increase in incoming diffuse light, which penetrates more deeply into the subcanopy, may increase the light levels in the subcanopy and the rates of photosynthesis by lower leaves. Others have suggested that increased radiation sharing because of varied leaf orientation within the canopy instead of increased beam penetration is the main cause of the increase in light use efficiency [Alton et al., 2007].

[5] Atmospheric aerosols may also influence ecosystems by reducing leaf temperature, which decreases respiration and heat stress on leaves. Field observations have drawn attention to the possibility that tropical forests are highly sensitive to temperature [Clark et al., 2003; Doughty and Goulden, 2008a; Feeley et al., 2007]. Field measurements of whole ecosystem CO$_2$ exchange in tropical forests made using the eddy covariance technique indicate that CO$_2$ uptake decreases markedly during warmer than average periods [Loescher et al., 2003; Goulden al., 2004]. These observations raise the possibility that aerosol induced cooling increases CO$_2$ uptake by tropical forests. Some recent studies have suggested that temperature may be as important as diffuse radiation in increasing CO$_2$ uptake during periods of increased optical depth [Min, 2005; Steiner and Chameides, 2005].

[6] The effect of biomass burning on surface radiation balances can be quantified using AERONET monitoring stations, which measure the increase in column light extinction relative to pure Rayleigh scattering [Holben et al., 1998]. Aerosol concentrations in the Amazon reached 30,000 particles per cubic centimeter in the dry season and an aerosol optical thickness of 3 at 500 nm [Andreae et al., 2004; Procopio et al., 2004], although the Tapajos region rarely exceeded AOT > 1.

[7] A previous study quantified the effects of biomass burning on the Amazon by comparing CO$_2$ fluxes from eddy covariance towers with aerosol optical depth in two regions of the Amazon forest [Oliveira et al., 2007]. Net ecosystem exchange (NEE) increased during periods of high optical depth associated with biomass burning. The authors could not determine if increased NEE was caused by an increase in subcanopy light or a decrease in canopy temperature.

[8] In this paper we investigated whether increased aerosol optical depth in an Amazonian forest increased NEE because of increased diffuse light or decreased temperature. We used measurements of aerosol optical depth, eddy covariance CO$_2$ flux, subcanopy PPFD, emitted long-wave radiation for canopy skin temperature, and subcanopy leaf-level photosynthesis to test three hypotheses:

Hypothesis 1: For a given global irradiance, diffuse light in the subcanopy increases during smoky, noncloudy periods compared to nonsmoky, noncloudy periods.

Hypothesis 2: For a given global irradiance, NEE increases during noncloudy, high AOT periods, because of increased irradiance of subcanopy shade leaves.

Hypothesis 3: Applying current understanding of factors governing NEE, the elevated NEE rates that are observed under smoky conditions cannot be explained solely by enhanced photosynthesis because of increased diffuse light.

2. Methods

2.1. Site

[9] This study was conducted at the Large-scale Biosphere-Atmosphere project (LBA-ECO) km–83 site (3.020833° S, 54.972221° W) in the Floresta National do Tapajos (FLONA). The FLONA Tapajos extends from 50 to 150 km south of Santarem, Para, Brazil, on the eastern side of the Tapajos River. The vegetation was closed tropical forest, with a canopy height of ∼40 m and scattered emergent trees up to 55 m. The FLONA was in the 27th percentile (∼2% to 3%) of Amazonian forests with respect to both annual precipitation and wet season length [Saleska et al., 2003]. The forest was on a broad, flat plateau [Goulden et al., 2004]. Soils were mainly yellow latosol clay (Haplic acrochord). An area extending 2 to 3 km east of the main km–83 eddy covariance tower was selectively logged in September 2001 as part of the larger Large-scale Biosphere-Atmosphere project (LBA-ECO) experiment, and its effect is described in previous papers [Miller et al., 2007; Figueira et al., 2008].

2.2. Tower Measurements

[10] We used the eddy covariance method [Wofsy et al., 1993; Baldocchi, 2003a] to measure the turbulent fluxes of CO$_2$ at 64 m above ground level. Tower equipment and data processing are explained in previous papers [da Rocha et al., 2004; Goulden et al., 2004; Miller et al., 2004]. Tower top meteorological measurements were recorded at 0.5 Hz. Incoming photosynthetically active photon flux density (PPFD) at 64 m was measured with silicon quantum sensors (LI-COR LI190; Lincoln, Neb.). We measured air temperature with ventilated thermisters (model 076B; Met One, Grants Pass, Ore; model 107, Campbell Scientific, Logan, Utah) and long-wave radiation emitted by the forest (CG2 net pyrgeometer; Kipp and Zonen, Delft, Holland). Below-canopy PPFD was measured with 8 quantum sensors (Apogee Instruments, Logan, Utah) that were horizontally mounted at ∼30 cm above the forest floor. The sensors were arrayed at 5 m intervals along two 15 m transects in a forest patch that was ∼20 to ∼40 m from the nearest logging-created gap and that was ∼40 m from the tower. This is an adequate sample size to get a representative sample of indirect shaded light [Reifsnyder et al., 1971].

2.3. Analysis of Half Hour Averaged Tower Data

[11] We calculated forest light curves using half hourly averaged eddy flux data from October–December 2002 and October–December 2003. We compared light curves on smoky days (at aerosol optical depths >0.5, >0.7, >1) versus nonsmoky days (AOT < 0.35). We compared half hour averaged gross ecosystem exchange (GEE), which was cal-
remote regions, aerosol optical depths at this distance are well correlated \cite{Andreae2004, Procopio2004}. AERONET has an algorithm to exclude cloud contamination of the AOT measurements \cite{Holben1998}. We screened the km–83 eddy flux data with three daily averaged optical depths thresholds: nonsmoky (AOT < 0.35), smoky (AOT > 0.5), and very smoky (AOT > 0.7). We used AERONET level 2.0 (cloud screened and recalibration of the CIMEL Sun photometer) daily averaged 500 nm aerosol optical depth from the nearby Belterra site (02°39'S and 54°58'W) (~40 km from our tower site) to determine AOT above our eddy covariance tower. We used data from October to December because these months had both high aerosol optical depth, relatively steady leaf area index (LAI), net ecosystem exchange (NEE) \cite{Doughty2008a}, and soil moisture \cite{Bruno2006}. We compared the measured decrease in downwelling shortwave irradiance to the predicted decrease in downwelling irradiance at different aerosol optical depths. The predicted decrease, calculated using a radiative transfer model modified for the Amazon, was based on digitized data from the study of Oliveira et al. \cite{2007}.

\section*{2.6. Analysis of Radiation Data}

\cite{2014} We used the tower top PPFD sensor to find 10 min, midday (~1100–1300 local time), nonsmoky (PPFD > 1200 \text{ \ mu mol m}^{-2} \text{ s}^{-1}), or cloudy intervals (PPFD < 800 \text{ \ mu mol m}^{-2} \text{ s}^{-1}) for nonsmoky (AOT < 0.35) days, smoky (AOT > 0.5) and very smoky days (AOT > 0.7). We compared irradiance between the tower top and subcanopy PPFD sensors to determine the effect of aerosol optical depth on subcanopy light. We removed all observations with a subcanopy PPFD greater than 10 \text{ \ mu mol m}^{-2} \text{ s}^{-1} to eliminate sunflecks and isolate the subcanopy shaded irradiance. We selected a threshold of 10 \text{ \ mu mol m}^{-2} \text{ s}^{-1}-based analyses of the distribution of subcanopy light which show a substantial decrease between light levels above 10 \text{ \ mu mol m}^{-2} \text{ s}^{-1} and much higher light levels (Figure 1). We tried several different thresholds to screen for light flecks, and all resulted in similar results. The PPFD sensors have the capability to measure changes of 1 \text{ \ mu mol m}^{-2} \text{ s}^{-1} in irradiance (Apogee Instruments, Logan, Utah, personal communication), and we programmed our data storage systems (Campbell CR23X) with sufficient bit size to record such a change. The light sensors are truly linear to complete darkness (Apogee Instruments, Logan, Utah, personal communication), and the observed sensor signals at night sensors were zero. We used midday periods (~1100–1300 local time) to avoid confounding zenith angle effects.

\cite{2015} We confirmed that the tower top and subcanopy sensors were synchronized by maximizing the covariance between the two data sets. We consider each 10 min period as a single sample for statistical analysis, in order to avoid pseudo replication.

\cite{2016} We calculated canopy temperature using a pyrgeometer to measure long-wave emission (L) (W m^{-2}). We calculated canopy skin temperature as

\begin{equation}
T = 
\left( \frac{L}{\varepsilon \sigma} \right)^{0.25},
\end{equation}

\begin{figure}[h]
\centering
\includegraphics[width=\textwidth]{figure1.png}
\caption{Seven ground level PPFD sensors averaged together during midday (1100–1300) noncloudy (PPFD > 1200 \text{ \ mu mol m}^{-2} \text{ s}^{-1}) smoky (AOT > 0.7) (gray line) (n = 13,395) and noncloudy (PPFD > 1200 \text{ \ mu mol m}^{-2} \text{ s}^{-1}) non-smoky periods (AOT < 0.35) (black line) (n = 34,651). Data are normalized histograms with bin widths of 5 \text{ \ mu mol m}^{-2} \text{ s}^{-1}.
\end{figure}

\section*{2.4. Analysis of 5 Min Averaged Tower Data}

\cite{2012} To increase our sample size and separate noncloudy and cloudy periods, we screened the data to find 5 min periods that were consistently noncloudy or cloudy and used these periods to calculate the eddy covariance fluxes. Eddy covariance fluxes are usually calculated using a half-hour averaging time to ensure that the low frequency contributions to the flux are captured \cite{Baldocchi2003b}. There are few continuously noncloudy or cloudy half-hour periods in the tropics. By calculating fluxes at 5 min intervals, we increased our sample size and reduced random sampling error \cite{Doughty2008a}. We screened the tower top PPFD data to identify intervals with either 10 cloudy minutes (PPFD < 800 \text{ \ mu mol m}^{-2} \text{ s}^{-1}) or 10 noncloudy minutes (PPFD > 1200 \text{ \ mu mol m}^{-2} \text{ s}^{-1}). We calculated the turbulent fluxes of CO\textsubscript{2} for the second 5 min interval to avoid transient effects that may be present in the first 5 min interval.
Table 1. AERONET Aerosol Optical Depth at 500 nm for the Belterra Site Between October–December 2002 and October–December 2003

<table>
<thead>
<tr>
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<th></th>
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</thead>
<tbody>
<tr>
<td>AOT &lt; 0.35</td>
<td>0.23</td>
<td>68</td>
<td>0.94</td>
<td>11%</td>
<td>8%</td>
</tr>
<tr>
<td>AOT &gt; 0.5</td>
<td>0.71</td>
<td>32</td>
<td>0.84</td>
<td>14%</td>
<td>13%</td>
</tr>
<tr>
<td>AOT &gt; 0.7</td>
<td>0.91</td>
<td>12</td>
<td>0.81</td>
<td>14%</td>
<td>13%</td>
</tr>
</tbody>
</table>

*We found averaged AOT and the number of days where AOT < 0.35, >0.5 and >0.7. We calculated expected relative irradiance for AOT < 0.35, >0.5 and >0.7 for the Belterra site following [Oliveira et al., 2007]. We compared the expected relative irradiance due to aerosols to the actual decrease in shortwave downwelling irradiance at our site.

2.7. Leaf Gas Exchange

[17] We measured leaf gas exchange from scaffold towers [Doughty and Goulden, 2008b; Doughty et al., 2006] at the LBA–ECO km–83 [Goulden et al., 2004] and km–67 [Saleska et al., 2003] field sites. Both sites were in closed-canopy evergreen forest in the FLONA Tapajós. The km–83 scaffold tower was located within 300 m of the main eddy flux tower. We measured the following subcanopy and understory species: *Eschweileria amazonica*, *Lecythis sp.*, *Chimarrhis turbinata*, *Faramea platyneura*, *Sclerolobium paraense*.

[18] We used a portable gas exchange system (LI 6400; Li-Cor Biosciences, Lincoln, Neb) in April and May 2002 to measure the rates of gas exchange at a range of light levels (PPFD 0, 50, 100, 200, 500, 1000, 1500, 1800 μmol m–2 s–1), with constant temperature (30°C) and ambient CO2. We grouped the data into understory (0–10 m) and subcanopy (10–25 m) leaves.

2.8. Canopy Radiation Transfer Model

[19] To model the partitioning of diffuse and direct radiation beneath the canopy top, we applied an adapted version of the shortwave narrowband model [Zender et al., 1997], which computes multistream fluxes using DISORT [Stamnes et al., 1988]. We used 10 canopy layers. We assumed a spherical leaf angle distribution, such that optical depth within each layer was 0.5 × (leaf + stem area index). We assumed leaves were the fundamental scattering element, and single-scatter albedo equaled leaf albedo. We applied the suggested aerosol optical properties for tropical forest from the study of Reid et al. [2005], extending 550 nm values to the full spectrum using a Mie solver.

[20] The model predictions of forest floor light did not match the observations: the model predicted a larger change between smoky and nonsmoky periods than was observed. Our PPFD sensors originally had a median value of 1.5 μmol m–2 s–1 during low AOT periods. This is lower than previous measured wet tropical forest subcanopy light which averaged between 5 and 50 μmol m–2 s–1 [Chazdon and Pearcy, 1991]. Our sensors darkened over several years because of downward drift in the signals associated with aging or the accumulation of dirt on the instrument surfaces [Doughty and Goulden, 2008b]. We calculated the average downward linear trend for all sensors and used these regressions to correct for the downward drift in the understory sensors. This increased median peak diffuse understory irradiance to 7.3 μmol m–2 s–1 at low AOTs (Figure 1) which is within the range found by others [Chazdon and Pearcy, 1991]. With these corrections, the model estimation of increased diffuse light because of smoke was close to our subcanopy sensors but still almost double our measurements. Therefore, to calculate the increase in diffuse light at each LAI, we divided the vertical profile model output by the model at each LAI by the ratio of the bottom of canopy light observed versus modeled, which was ~2.

[21] We calculated the percentage LAI shaded at each LAI and multiplied this by the increase in diffuse light with the quantum yield (carbon fixed per quanta light) of photosynthesis from 0 to 50 μmol m–2 s–1. We determined the change in net canopy photosynthesis due to increased shaded light as

\[ \sum_{i=1}^{5} \Delta D(LAI_i) \times S(LAI_i) \times Q(LAI_i), \]

where \( \Delta D \) is the increase in subcanopy light at each LAI, \( S \) is percentage LAI shaded at each LAI, and \( Q \) is the quantum yield for leaves at each LAI. We assumed an LAI of 6 [Doughty and Goulden, 2008b].

3. Results

3.1. Effect of Aerosol Optical Depth on Irradiance

[22] There were 68 days at Belterra when the daily averaged 500 nm AERONET aerosol optical depth was below 0.35, 32 days when it was above 0.5, and 12 days when it was above 0.7, during October–December 2002 and October–December 2003. The average AOT was 0.23 at AOT < 0.35, 0.71 for AOT > 0.5, and 0.91 for AOT > 0.7. We compared measured decreases in irradiance due to smoke for the Tapajos region to modeled decreases [Oliveira et al., 2007] to ensure that the decrease in irradiance during smoky periods was due to the smoke. Modeled irradiance predicted a 11% decline between AOT 0.23 and AOT 0.72 and a 14% decline between AOT 0.23 and AOT 0.91 at zenith angles between 20° and 30° [Oliveira et al., 2007]. Averaged downwelling shortwave irradiance at the km–83 field site decreased during noncloudy periods (defined as >1200 μmol m–2 s–1 irradiance) between 1100 and 1300 local time by 8% between averaged AOT = 0.23 and averaged AOT = 0.72 and decreased by 13% between averaged AOT = 0.23 and averaged AOT = 0.72.
0.91 (Table 1). Measured changes in irradiance, due to smoke, closely matched modeled changes in irradiance.

The occurrence of periods with irradiance above 1800 μmol m⁻² s⁻¹ decreased substantially during smoky times (AOT > 0.5) compared with nonsmoky intervals (AOT < 0.35) (Figure 2). However, contrary to our expectations, irradiance increased slightly during cloudy periods (irradiance between 600–800 μmol m⁻² s⁻¹) that were also smoky (AOT > 0.5) relative to nonsmoky intervals (AOT < 0.35).

3.2. Effect of Aerosol Optical Depth on Half-Hour Averaged Canopy Fluxes

We compared half-hour averaged gross ecosystem exchange (GEE) between October to December at various optical depths (AOT > 0.5, AOT > 0.7, and AOT > 1) (Figure 3). There were no significant differences between light-saturated GEE (PPFD > 1000 μmol m⁻² s⁻¹) during nonsmoky periods (AOT < 0.35) and smoky periods (AOT > 0.5). However, there was a significant increase in light-saturated GEE at AOT > 1 (P < 0.05) and the increase in GEE at AOT > 0.7 was large and nearly significant (P = 0.08). The mean light-saturated GEE (PPFD > 1000 μmol m⁻² s⁻¹) ± standard deviation at AOT < 0.35 was −19.8 ± 7.1, −19.8 ± 8.0 at AOT > 0.5, −22 ± 6.5 at AOT > 0.7, and −25.8 ± 5.1 at AOT > 1.

3.3. Effect of Aerosol Optical Depth on 5 Min Averaged Canopy Fluxes

We calculated 5 min averaged NEE for smoky and nonsmoky periods. CO₂ uptake increased significantly (P > 0.05) during smoky, noncloudy periods (AOT > 0.5) and during (P > 0.001) very smoky, noncloudy periods (AOT > 0.7) compared to nonsmoky (AOT < 0.35), noncloudy periods (Figure 4). CO₂ uptake at AOT > 0.7 was significantly greater (P > 0.05) than CO₂ uptake at AOT > 0.5. There were no significant differences between smoky, cloudy, and nonsmoky, cloudy periods, although there was a slight decrease in mean CO₂ uptake during smoky, cloudy periods.

3.4. Effect of Aerosol Optical Depth on Subcanopy Shaded Light

Subcanopy shaded light increased slightly during smoky, noncloudy periods versus nonsmoky, noncloudy periods (AOT > 0.5 and AOT > 0.7), cloudy (PPFD < 800 μmol m⁻² s⁻¹) nonsmoky (AOT < 0.35), and nonsmoky (PPFD > 1200 μmol m⁻² s⁻¹) nonsmoky periods (AOT < 0.35). All 5 min flux calculations were preceded by a 5 min period with identical conditions to avoid edge effects. The errorbars are ±SE. *** indicates significance at the P < 0.001. * indicates significance at the P < 0.05.
periods (Figure 1). Light levels were significantly greater ($P < 0.05$) during smoky and very smoky periods ($AOT > 0.5$, $> 0.7$) than nonsmoky periods ($AOT < 0.35$) during midday (1100–1300 local time) noncloudy periods (PPFD $> 1200 \mu$mol m$^{-2}$ s$^{-1}$) (Figure 5). There were no significant differences in subcanopy shaded light between midday (1100–1300 local time), cloudy (PPFD $> 400$ and PPFD $< 800 \mu$mol m$^{-2}$ s$^{-1}$), nonsmoky ($AOT < 0.35$), and midday, cloudy, smoky ($AOT > 0.5$) periods. There was a significant decrease in subcanopy shaded light during cloudy, very smoky $AOT > 0.7$ periods versus cloudy, nonsmoky $AOT < 0.35$ periods.

[27] The ratio of irradiance reaching the subcanopy is significantly higher during cloudy intervals than noncloudy intervals (0.035 cloudy at $AOT < 0.35$ versus 0.01 noncloudy at $AOT < 0.35$). However, the actual subcanopy PPFD levels (after removing sunflecks) are only slightly higher during cloudy versus noncloudy intervals ($21.5 \mu$mol m$^{-2}$ s$^{-1}$ cloudy at $AOT < 0.35$ versus $19.7 \mu$mol m$^{-2}$ s$^{-1}$ noncloudy at $AOT < 0.35$).

[28] Subcanopy shaded light during noncloudy periods increased $2.9 \mu$mol m$^{-2}$ s$^{-1}$ at $AOT > 0.5$ (average optical depth 0.72) and $4.1 \mu$mol m$^{-2}$ s$^{-1}$ at $AOT > 0.7$ (average optical depth 0.91). When plotted with an increase of $0 \mu$mol m$^{-2}$ s$^{-1}$ at $AOT < 0.35$ (average optical depth 0.23), there was a slope of $6.1 \mu$mol m$^{-2}$ s$^{-1}$ $AOT^{-1}$, which means that for every increase in $AOT$ of 1, subcanopy shaded PPFD increased $\sim 6.1 \mu$mol m$^{-2}$ s$^{-1}$ (Figure 6), an increase of $\sim 30\%$ assuming a mean subcanopy noncloudy PPFD of $20 \mu$mol m$^{-2}$ s$^{-1}$.

### 3.5. Effect of Aerosol Optical Depth on Canopy Temperature

[29] We quantified the decline in canopy temperature due to decreased irradiance from smoke. We subtracted air temperature from canopy temperature because air temperature, a controlling factor in canopy temperature, often shows variability not related to changing irradiance. During very smoky periods ($AOT > 0.7$), downwelling shortwave irradiance decreased by 123 W m$^{-2}$ relative to nonsmoky periods ($AOT < 0.35$) (Table 1: 123 W m$^{-2}$ $= 13\%$ measured decrease in solar downwelling radiation). Canopy skin temperatures cooled relative to air temperatures by 0.41°C during very smoky periods ($AOT > 0.7$) versus nonsmoky periods ($AOT < 0.35$). During smoky periods ($AOT > 0.5$), downwelling shortwave irradiance decreased by 79 W m$^{-2}$ relative to nonsmoky periods ($AOT < 0.35$). Canopy skin temperatures cooled relative to air temperatures by 0.26°C in smoky periods ($AOT > 0.5$) versus nonsmoky periods ($AOT < 0.35$). A linear regression had a slope of $\sim 300$ W m$^{-2}$ C$^{-1}$, meaning that canopy temperature decreases 1°C relative to air temperature for every 300 W m$^{-2}$ decrease in downwelling solar radiation (Figure 7).

![Figure 5](image5.png)

**Figure 5.** Noontime data (1100–1300) for (a) tower top PPFD during periods of cloudy (PPFD $< 800 \mu$mol m$^{-2}$ s$^{-1}$) smoky ($AOT > 0.5$ and $AOT > 0.7$), cloudy (PPFD $< 800 \mu$mol m$^{-2}$ s$^{-1}$) nonsmoky ($AOT < 0.35$), noncloudy (PPFD $> 1200 \mu$mol m$^{-2}$ s$^{-1}$) smoky ($AOT > 0.5$ and $AOT > 0.7$), and noncloudy (PPFD $> 1200 \mu$mol m$^{-2}$ s$^{-1}$) nonsmoky periods ($AOT < 0.35$). (b) The amount of PPFD reaching the subcanopy during these periods. The errorbars are 95% confidence intervals. (c) The ratio between PPFD at the top of the canopy and diffuse light (nonsunflecks) intercepted by seven ground level subcanopy PPFD sensors during the same intervals.

![Figure 6](image6.png)

**Figure 6.** Increase in subcanopy PPFD measured by seven subcanopy PPFD sensors after removing all light flecks during noncloudy (tower top PPFD $> 1200 \mu$mol m$^{-2}$ s$^{-1}$) noontime (1100–1300 local time) periods where averaged daily averages of $AOT > 0.5$ and $> 0.7$ are subtracted from daily averages when $AOT < 0.35$. This was plotted against mean $AOT$ for $AOT > 0.5$, $> 0.7$, and $< 0.35$. There are no error bars because we subtract means with different sample sizes.
Figure 7. Mean downwelling shortwave irradiance during noontime (1100–1300 local time) and AOT > 0.5 and AOT > 0.7 subtracted against mean noontime downwelling shortwave irradiance at AOT < 0.35 from October–December 2002 and 2003. These data are plotted against average canopy temperature measured with a longwave pyrgeometer subtracted from average air temperature for the same periods of time. There are no error bars because we subtract means with different sample sizes. These data are plotted through the origin, which is the mean shortwave irradiance and mean canopy minus air temperature at AOT < 0.35.

[30] Air temperatures cooled slightly, by 0.15°C, between very smoky (AOT > 0.7) and nonsmoky (AOT < 0.35) periods. Surprisingly, air temperatures warmed slightly between smoky (AOT > 0.5) and nonsmoky (AOT < 0.35) periods. Even though the air was slightly warmer at smoky (AOT > 0.5) than nonsmoky (AOT < 0.35) periods, the overall canopy temperature still cooled because of decreased solar irradiance.

3.6. Whole Canopy CO$_2$ Exchange as a Function of Skin Temperature

[31] Previous analyses of the relationship between irradiance and CO$_2$ exchange at the km-83 site indicated that whole canopy photosynthesis saturates at a PPFD of 1000 μmol m$^{-2}$ s$^{-1}$ [Goulden et al., 2004]. We therefore analyzed the rates of GEE during periods with PPFD above 1000 μmol m$^{-2}$ s$^{-1}$ to determine the sensitivity of light-saturated CO$_2$ exchange to canopy skin temperature (Figure 8). GEE was relatively insensitive to changes in canopy skin temperature below 28°C and highly sensitive to changes in canopy skin temperature above 28°C. We analyzed a subset of the data using 5 min averaged NEE and canopy temperature from midday (1000–1400 local time), October–December 2002 and October–December 2003, noncloudy (PPFD > 1200 μmol m$^{-2}$ s$^{-1}$), nonsmoky (AOT < 0.35) periods to ensure that the temperature effect was due to temperature and not clouds or aerosols (Figure 8). The forest had a similar response of CO$_2$ uptake to temperature in both data sets (a ~2 μmol m$^{-2}$ s$^{-1}$ decrease in canopy CO$_2$ uptake for every 1°C increase in temperature above 28°C, a ~13% decrease).

3.7. Leaf Photosynthesis

[32] We made 25 light photosynthesis curves on leaves below 10 m and 42 light photosynthesis curves on leaves between 10 and 25 m from the ground between April and May 2002 to quantify the effect of increased shaded light on subcanopy leaf photosynthesis. Leaves below 10 m became light saturated at ~400 μmol m$^{-2}$ s$^{-1}$, with a maximum photosynthetic capacity ($A_{\text{max}}$) of 3 μmol m$^{-2}$ s$^{-1}$ (Figure 9a). Leaves between 10 and 25 m became light saturated at ~600 μmol m$^{-2}$ s$^{-1}$, with an $A_{\text{max}}$ of ~5 μmol m$^{-2}$ s$^{-1}$. We linearly regressed photosynthesis and irradiance at light levels of 0 and 50 μmol m$^{-2}$ s$^{-1}$ and found a quantum yield of 0.039 ± 0.57 at 0–10 m and a quantum yield of 0.039 ± 0.83 at 10–25 m (Figure 9b). These quantum yields are less than those typically reported for C$_3$ dicots (0.052 mol CO$_2$ E$^{-1}$; [Ehleringer and Peary, 1983]). This difference may be due in part to our calculation of quantum yield based on the incident light, as opposed to the true quantum yield, which is based on absorbed light.

4. Discussion

[33] We focused our study on the 3 months during the dry season, when AOT was highest (October to December), rather than the whole year, as done by Oliveira et al. [2006]. Our focus on this period reduced the confounding influence of seasonal changes in LAI [Doughty and Goulden, 2008a], soil moisture [Bruno et al., 2006], and specific humidity [da Rocha et al., 2004]. We found no significant differences between smoky and nonsmoky periods for soil moisture (volumetric soil moisture at 20 cm AOT > 0.7 = 0.390 m$^3$ m$^{-3}$ above 28°C, a ~13% decrease).
versus $\text{AOT} < 0.35 = 0.394 \text{ m}^2 \text{ m}^{-3}$ for 1100–1300 local time, October–December 2002 and October–December 2003), $\text{LAI (AOT} > 0.7 = 6.26 \text{ m}^2 \text{ m}^{-2}; \text{AOT} < 0.35 = 6.23 \text{ m}^2 \text{ m}^{-2})$ [Doughty and Goulden, 2008b], or specific humidity (AOT > 0.7 = 27.90 mmol mol$^{-1}$; AOT < 0.35 = 27.63 mmol mol$^{-1}$).

[34] We found a significant difference in light-saturated half-hour GEE between AOT > 1 and AOT < 0.35, and a large, nearly significant ($P = 0.08$), difference between AOT > 0.7 and AOT < 0.35 (Figure 3). There was no difference in mean light-saturated GEE between AOT > 0.5 and AOT < 0.35. Half-hour averaged GEE often blend non-cloudy and cloudy periods which reduce the measured effect of smoky periods because cloudy, smoky periods do not increase GEE like non-cloudy, smoky periods.

[35] Five minute averaged NEE increased significantly ($P < 0.001$) between very smoky, non-cloudy and non-smoky, non-cloudy periods. NEE at AOT > 0.7 was substantially greater than NEE at AOT > 0.5, and both were greater than AOT < 0.35 (Figure 4). There were no significant changes in NEE between cloudy, smoky periods (AOT > 0.7) and cloudy, non-smoky periods (AOT < 0.35). However, despite having higher tower top light levels during smoky, cloudy periods, NEE was slightly lower in smoky periods than non-smoky periods, possibly due to decreased subcanopy shaded light at AOT < 0.35 versus at AOT > 0.7.

[36] There were small but significant increases in subcanopy diffuse irradiance during noncloudy, smoky (AOT > 0.5) and noncloudy, very smoky periods (AOT > 0.7) compared to noncloudy, nonsmoky periods (AOT < 0.35) (Figure 5). We multiplied the modeled increase in diffuse light (scaled to match the bottom of canopy PPFD measurements) with increased AOT at each LAI by the quantum yield at each LAI to get a net canopy scale noncloudy midday estimated increase in CO$_2$ uptake of 3.8 $\mu$mol m$^{-2}$s$^{-1}$.

[37] Decreased irradiance with smoke cooled canopy skin temperature by 0.4°C relative to air temperature (Figure 7), which could impact NEE during peak midday temperatures. The total cooling ($\Delta$ air temperature + $\Delta$ canopy temperature) between canopy temperature at AOT > 0.7 and AOT < 0.35 was ~0.55°C. There was a linear drop ($r^2 = 0.88$) in CO$_2$ uptake at canopy skin temperatures above 28°C with a slope of ~1.93 $\mu$mol m$^{-2}$ s$^{-1}$ C$^{-1}$ (Figure 8). A decrease in temperature of 0.55°C would increase CO$_2$ uptake by ~1.1 $\mu$mol m$^{-2}$ s$^{-1}$ at air temperatures above 28°C. The combined effects of temperature and diffuse light should result in an increase in CO$_2$ uptake of ~4.9 $\mu$mol m$^{-2}$ s$^{-1}$, with decreased tem-

![Figure 9. (a) Photosynthesis $\mu$mol m$^{-2}$ s$^{-1}$ ± SD at various light levels (30°C and ambient CO$_2$) for species at the km-83 and km-67 field sites below 10 m and between 10 and 25 m. Measurements were made on canopy platform towers between April and May 2002. (b) Quantum yield in the subcanopy as determined by the increase in photosynthesis from 0 to 50 $\mu$mol m$^{-2}$ s$^{-1}$ for species below 10 m and species between 10 and 25 m.](image-url)
temperature accounting for 22% of the effect and increased diffus radiation to the subcanopy accounting for 78% of the effect.

[38] Our site has an average daytime noncloudy NEE of $-12.5 \text{ mmol m}^{-2} \text{ s}^{-1}$ (which is an underestimate of real NEE due to the 5 min averaging period) at AOT $< 0.35$, $-14.4 \text{ mmol m}^{-2} \text{ s}^{-1}$ at AOT $> 0.5$, and $-17.6 \text{ mmol m}^{-2} \text{ s}^{-1}$ at AOT $> 0.7$. The combined effects of temperature and diffuse light are $-4.9 \text{ mmol m}^{-2} \text{ s}^{-1}$, and the difference in CO$_2$ fluxes between AOT $< 0.35$ and AOT $> 0.7$ is $-5.1 \text{ mmol m}^{-2} \text{ s}^{-1}$. CO$_2$ uptake could also be affected through increased relative humidity that increased with decreased air temperature during smoky periods. A model based on the Ball Berry stomatal conductance equation showed that the effect of increased humidity on CO$_2$ uptake is quite small ($\Delta < 1\%$) under small changes in relative humidity (a 0.15°C decrease in air temperature will increase humidity by 0.73% which will change CO$_2$ uptake by $0.2 \text{ mmol m}^{-2} \text{ s}^{-1}$) [Collatz et al., 1991].

[39] In this study, changes in seasonal LAI, soil moisture, and specific humidity were not significantly different between periods of high and low AOT, and the effect of changing relative humidity between these periods on CO$_2$ uptake is likely small. The predicted combined effect of temperature and diffuse light closely matches experimental observations. We therefore feel confident in our prediction that the increase in CO$_2$ uptake during smoky periods is real and about 20% due to a reduction of temperatures and about 80% due to increased subcanopy diffuse radiation. We calculate that during the period of study, aerosols may have increased average daytime noncloudy NEE at our site by $-2.7\%$ over an entire year ($321 \text{ days AOT} < 0.35 \times 12.5 \text{ mmol m}^{-2} \text{ s}^{-1} + 32 \text{ days AOT} > 0.5 \times 14.4 \text{ mmol m}^{-2} \text{ s}^{-1} + 12 \text{ days AOT} > 0.7 \times 17.6 \text{ mmol m}^{-2} \text{ s}^{-1} = 365 \text{ days AOT} < 0.35 \times 12.5 \text{ mmol m}^{-2} \text{ s}^{-1}$). 

[40] It is difficult to know whether this proportion between temperature and diffuse light is applicable to forests worldwide or limited to tropical forests. Current evidence suggests small perturbations in temperatures have a large effect on tropical forests [Doughty and Goulden, 2008a; Clark et al., 2003] possibly because annual and seasonal temperature changes in tropical forests are small. Temperate forests, with their larger variations in seasonal and annual temperatures, are likely to have smaller increases in NPP due to decreased temperatures during periods of high AOT than tropical forests. The effect of AOT on surface air temperatures at our site was much smaller than in previous studies [Schafer et al., 2002], and temperature effects may have a larger effect on NPP during other high aerosol events that more strongly affect surface air temperatures.

[41] Overall, aerosols increased CO$_2$ uptake at our site mainly due to increased subcanopy light (80%) and partially due to lower canopy temperatures (20%). Future studies are necessary to determine whether this increased CO$_2$ uptake would lead to increased tropical forest tree growth or increased respiration.

[42] Acknowledgments. We thank Helber Freitas, Michela Figueira, Augusto Maia, and Albert de Sousa for field help; Scott Miller and Humberto da Rocha for help with the tower measurements; UFPA in Santarem and Belem for help identifying species; Jim Randerson for comments; Paulo Artaxo and Brent Holben for the AERONET data; and the LBA-ECO Santarem field office for support. This work was funded by NASA, as a component of LBA-ECO and an Earth System Science Fellowship, and NSF-0738369.

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