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Confirmation of progressive plate motion during the Midcontinent Rift’s early magmatic stage from the Osler Volcanic Group, Ontario, Canada

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Abstract

As the supercontinent Rodinia was assembling ca. 1.1 billion years ago, there was extensive magmatism on at least five Proterozoic continents including the development of the North American Midcontinent Rift. New paleomagnetic data from 84 lava flows of the Osler Volcanic Group of the Midcontinent Rift reveal that there was a significant and progressive decrease in inclination between the initiation of extrusive volcanism in the region (ca. 1110 Ma) and ca. 1105 ± 2 Ma (during the “early stage” of rift development). Paleomagnetic poles can be calculated for the lower portion of the reversed Osler Volcanic Group (40.9° N, 218.6° E, A95 = 4.8°, N = 30) and the upper portion of the reversed Osler Volcanic Group (42.5° N, 201.6° E, A95 = 3.7°, N = 59; this pole can be assigned the age of ca. 1105 ± 2 Ma). This result is a positive test of the hypothesis that there was significant plate motion during the early stage of rift development. In addition to being a time of widespread volcanism on Laurentia and other continents, this interval of the late Mesoproterozoic was characterized by rapid paleogeographic change.

1. Introduction

Despite being active for more than 20 million years [Davis and Green, 1997] and resulting in the thinning of prerift crust to less than 10 km [Cannon, 1992], the 1.1 Ga Midcontinent Rift failed to dismember Laurentia (cratonic North America). This failure resulted in the preservation of a thick record of rift-related volcanic and sedimentary rocks that gives geoscientists insight into the development of this ancient rift. Most models for the development of the Midcontinent Rift attribute its origin to the upwelling and decompression melting of a mantle plume [Shirey, 1997]. On the basis of the great volume of generated magma and interpretation of geochemical data, it is argued that the early stage plateau flood basalts of the rift (ca. 1110–1105 Ma) and the main stage volcanics that erupted into the central basin (ca. 1100–1095 Ma) were both dominated by plume-sourced melts. This deep-plume origin for the rift needs to be considered in conjunction with paleogeographic change that has been inferred to have been ongoing throughout rift development. Fully constraining this paleogeographic change is essential for understanding rift development and for constraining late Mesoproterozoic paleogeographic reconstructions given the centrality of Laurentia’s apparent polar wander path to such efforts.

It has long been noted that there is a significant difference in paleomagnetic inclination between the steep (dominantly reversed polarity) magnetizations of the oldest volcanics and intrusives of the Midcontinent Rift and the shallower (dominantly normal polarity) magnetizations from the younger main stage volcanics and intrusives [Halls and Pesonen, 1982]. This inclination change has been interpreted either as resulting from rapid plate motion [Robertson and Fahrig, 1971; Davis and Green, 1997] or as being the result of large nondipolar contributions to the late Mesoproterozoic geomagnetic field that led to asymmetry across reversals [Pesonen and Nevanlinna, 1981]. The interpretation of the record as recording stepwise inclination change across reversals associated with a significant sustained departure from a geocentric axial dipole (GAD) dominated field was challenged by the observation of a progressive decrease in paleomagnetic inclination across multiple geomagnetic reversals up through the succession of Midcontinent Rift lavas at Mamainse Point, Ontario [Swanson-Hysell et al., 2009, 2014]. This progressive decrease in inclination leads to the interpretation of multiple symmetric reversals when the data are considered in stratigraphic context.
and these results have been used to support the hypothesis that rapid paleogeographic change was ongoing during Midcontinent Rift magmatism [Buchan, 2013; Swanson-Hysell et al., 2014]. To date, Mamainse Point is the only succession where a progressive decrease in inclination through an exposure of rift stratigraphy has been reported. At Mamainse Point, much of the decrease in paleomagnetic inclination occurs within the lowermost reversed polarity portion of the stratigraphy that erupted during the early magmatic stage of rift development [Swanson-Hysell et al., 2014]. This result sets up the prediction that other localities in the rift that span the same period of time should also record such a decrease. This work tests that hypothesis with new paleomagnetic data developed in stratigraphic context from the Osler Volcanic Group.

2. Geology and Context of the Osler Volcanic Group

2.1. Osler Volcanic Group Lithologies

The Osler Volcanic Group overlies the epicontinental sediments of the Mesoproterozoic Sibley Group. The lowest 100 m of the Osler Volcanic Group contains rift-related sandstones and conglomerates [Hollings et al., 2007a], which are followed by a succession of relatively continuous tholeiitic basalt flows. We studied these flows along the east shore of Simpson Island in the Lake Superior Archipelago. This location has well-preserved basalt flows and exposure of a significant thickness of the Osler Volcanic Group stratigraphy. This location occurs at a great enough distance (>10 km) from the intrusive St. Ignace Island Complex so that the magnetizations of the flows are apparently unaffected by its emplacement. The St. Ignace Island Complex is dominated by felsic lithologies and was emplaced into the mafic flows of the Osler Volcanic Group near the end of Midcontinent Rift volcanism [Hollings et al., 2007b].

Along the east shore of Simpson Island, where we measured stratigraphic sections (Figure 1), the stratigraphic succession is dominated by basaltic lava flows with typical thicknesses of ca. 5 m. For the 105 flows in the measured stratigraphic sections (that represent a subset of the total stratigraphy) where there is sufficient exposure of the flow base, interior and top to determine thickness, the median flow thickness is 4.9 m with a first quartile thickness of 2.0 m and third quartile thickness of 9.8 m (details are provided in the supporting information). Minor interflow siltstone and conglomerate occur between some flows and provide constraints on paleohorizontal that were utilized for structural corrections of paleomagnetic data (Figure 1). The red-brown siltstone beds are generally centimeters to tens of centimeters thick while the conglomerate beds can be up to several meters in thickness. Clasts in the conglomerate beds are basaltic in composition and are primarily pebble sized, but can range up to 10 cm in diameter. Within the stratigraphic sections, individual flows can be distinguished by their characteristic transition in texture moving up through the flow: pipe vesicles at the base of flows, to massive basalt, to massive basalt with some amygdules, to a highly amygdaloidal texture at the top of the flows. There is well-preserved pahoehoe flow banding on some flow tops. Using our measurements of bedding orientation, we estimate that slightly over 3000 m of the Osler Volcanic Group is exposed on Simpson Island (Figure 1).

2.2. Angular Unconformity and Associated Polarity Change

Southwest of Simpson Island in the Nipigon Strait region there is exposure on Puff Island containing an angular unconformity marked by a conglomerate that separates underlying flows of reversed magnetic polarity from overlying flows of normal magnetic polarity [Halls, 1974]. This unconformity is stratigraphically higher than the top of the exposure at Simpson Island where all studied flows have reversed magnetic polarity. Only ca. 110 stratigraphic meters of normally magnetized flows are exposed south (i.e., stratigraphically above) of the unconformity on Puff Island before the sequence is covered by Lake Superior. This unconformity has been interpreted to be due to a period of quiescence of local volcanism [Halls, 1974]. Halls [1974] was the first to suggest that such quiescence may have been widespread throughout the Midcontinent Rift—an idea that is now incorporated into models of rift development that utilize interpretations of the distribution of U-Pb dates and has been termed the “latent stage” [Vervoort et al., 2007]. As occurs in sequences across the Midcontinent Rift, the lower reversed flows studied paleomagnetically by Halls [1974] have a steeper inclination, in an absolute sense, than the younger normal flows above the unconformity. The missing time evidenced by the angular unconformity likely correlates with the missing time inferred from radiometric dates on units within the North Shore Volcanic Group and the Powder Mill Group [Davis and Green, 1997; Zartman et al., 1997]. In these rift successions, the stratigraphic intervals where time is
inferred to be missing are associated with a switch from reversed to normal magnetic polarity. A new U-Pb date from Mamainse Point (1100.36 ± 0.25 Ma from the upper reversed polarity zone) demonstrates that, in contrast to successions with missing time or condensed stratigraphy, the sequence there is relatively complete thereby preserving additional geomagnetic reversals [Swanson-Hysell et al., 2014].

The paleomagnetic data that Halls [1974] developed from Osler Volcanic Group flows were in relatively close stratigraphic proximity below and above the unconformity (as the goal of the study was to confirm the presence of a geomagnetic reversal that had been inferred from aeromagnetic data) [Halls, 1972]. The aeromagnetic data demonstrated that the remainder of the flows below the unconformity (i.e., to the north) and stratigraphically below those studied by Halls [1974] are also of reversed magnetic polarity. On the basis of polarity and the geochronology discussed below, the interval has been correlated with the “early stage” of Midcontinent Rift development and the other basal sequences of lava flows within the rift. These early stage flows are interpreted to be plateau lavas that erupted over a broad geographic area prior to significant development of the main central rift graben that underlies present day Lake Superior [Cannon, 1992]. Due to being limited to the portion of the succession in close stratigraphic proximity to the angular unconformity, existing data have not permitted evaluation of whether or not there is a progressive change in paleomagnetic inclination through the reversed polarity flows as observed in the “early stage” volcanics at Mamainse Point [Swanson-Hysell et al., 2009].

2.3. Age Constraints on the Osler Volcanic Group

In some locales, a quartz-feldspar, porphyritic felsic unit occurs near the base of the Osler Volcanic Group for which a U-Pb zircon date of 1107.5 ± 4 Ma has been reported [Davis and Sutcliffe, 1985]. This date was
obtained from an outcrop of felsic porphyry on Black Bay Peninsula, ca. 40 km to the west of Simpson Island [Davis and Sutcliffe, 1985]. This unit was tentatively interpreted as extrusive [Davis and Sutcliffe, 1985; Lightfoot et al., 1991] and the date has been interpreted as constraining the time at which Osler Group volcanism commenced. New observations made of a quartz-feldspar porphyry unit on Simpson Island, mapped as equivalent to the unit from which the date was obtained [Giguere, 1975], provide additional evidence for the inference made by Giguere [1975] that this unit is actually intrusive. On Simpson Island, the basal sedimentary units of the Osler Volcanic Group are overlain by the quartz-feldspar porphyry which itself underlies the basalt flows (Figure 1). A thin (1–2 mm thick) veneer of basalt is variably present overlying the porphyry. The basalt veneer displays pahoehoe flow banding, which is unlikely to have developed if the flow was originally this thin, implying that the felsic unit intruded into the basalt, cutting into an originally thicker flow. An additional observation is that there are protrusions of porphyry surrounded by host basalt, providing further support for an intrusive relationship. If the Simpson Island intrusion is indeed equivalent to the Black Bay Peninsula unit, this evidence suggests that the 1107 ± 2 Ma date is a minimum age for the eruption of the first Osler basalt flows, rather than an absolute age for that point in the Osler Volcanic Group stratigraphic succession.

A sequence of quartz-feldspar phyric rhyolite flows occurs near the top of the magnetically reversed portion of the Osler Volcanic Group at Agate Point [Davis and Sutcliffe, 1985] (Figure 3; stratigraphically higher than the highest flow on Simpson Island). Davis and Green [1997] obtained a U-Pb zircon date from the Agate Point Rhyolite of 1105 ± 2 Ma which, if the extrusive interpretation is correct, is a robust age for that point in the Osler Volcanic Group stratigraphy.

3. Methods

We collected oriented samples for magnetic laboratory measurements during the course of measuring stratigraphic sections (Figure 1). Each site consists of an individual lava flow from which we collected 6–10 small (2 cm diameter) rock cores with a hand-held drill. These small core samples were oriented with a magnetic and sun compass, when possible, such that their spatial orientation is known. To minimize the visual impact of collecting these small cores along the pristine Lake Superior shoreline, we knocked out the portion of the outcrop from which they were collected.

At the Institute for Rock Magnetism, specimens were prepared from the samples and subjected to progressive thermal or alternating field (AF) demagnetization (Figure 2a). Initial results on sister specimens demonstrated the simplicity of the magnetizations and the similarity between results obtained through thermal...
and AF demagnetization (Figure 2a). Low-temperature remanence experiments were run on representative samples and loss of remanence across the Verwey transition demonstrates that the magnetic mineralogy is dominated by low-titanium magnetite (e.g., Figure 2b). These results revealed the ability of AF demagnetization to isolate the characteristic remanent magnetization (ChRM) held by magnetite with relatively small and variably present overprints being effectively removed by low-field AF steps. Given these results, the majority of flows were studied with AF demagnetization alone. Line fits were made to the demagnetization data using principal component analysis (Kirschvink, 1980).

4. Results and Discussion

Flow means from the data generated for 84 flows of the Osler Volcanic Group are summarized in Figures 1 and 3 and in a table in the supporting information. To consider whether the flows of the Osler Volcanic Group record progressive paleogeographic change, we take the approach of grouping and comparing data from the lower third of the sequence (0–1041 m; 30 flows), the middle third of the sequence (1041–2083 m; 20 flows), and the upper third of the sequence (2083–3124 m; 34 flows). To test whether these subsets of the data could have been drawn from a common mean, we apply both the Watson $V_w$ test with Monte Carlo simulation [Watson, 1983] and the bootstrap test for a common mean [Tauxe, 2010]. Full details associated with these statistical tests are provided in the supporting information. The results from these common mean statistical tests show that the directions from the lower third of the sequence cannot be distinguished from those from the middle third, nor can the directions from the upper third of the sequence be distinguished from those in the middle third. In contrast, directions from the lower third of the sequence can be distinguished from those of the upper third at the 95% confidence level. This result can be seen visually in Figure 3 as the $A_{95}$ ellipses associated with the directional means and the $A_{95}$ ellipses associated with the pole means do not overlap. The statistically significant difference between the populations of flow means in the lower and upper third of the stratigraphy, combined with the result that the middle third data have a mean that is an intermediate direction between the lower and upper means, supports the hypothesis that progressive plate motion was ongoing throughout the eruption of the Osler Volcanic Group with Laurentia moving to lower latitudes (see reconstruction Figure 3). The paleomagnetic poles calculated and used in the paleogeographic reconstruction in Figure 3 are: lower third Simpson Island Osler Group (40.9° N, 218.6° E, $A_{95} = 4.8°$, $N = 30$); middle third Simpson Island Osler Group (42.7° N, 211.3° E, $A_{95} = 8.2°$, $N = 20$); and the upper portion of the reversed Osler Group (41.6° N, 205.4° E, $A_{95} = 4.8°$, $N = 34$). Stratigraphic subgroups of the Simpson Island data can be made in many different ways than this approach of dividing the stratigraphic sequence into thirds. For example, comparing the 17 flows in the lowermost 500 m against the 17 flows in the uppermost 500 m demonstrates that those populations are dramatically different such that the bootstrap test for a common mean shows their x, y, and z components to be distinct at the 99% confidence level (see supporting information for details). Comparing the 41 flows in the lower half of the stratigraphic succession to the 43 flows in the upper half of the stratigraphic succession also reveals a statistically significant, but relatively small, difference between the populations (see supporting information for details). We focus on the lower, middle, and upper third grouping in our analysis making the judgment that such an analysis strikes the balance between considering the possibility of change through the stratigraphy while binning enough data over thick enough intervals to not to be significantly biased by underaveraging secular variation. In this grouping, each bin contains $>1000$ m of relatively thin pahoehoe basalt flows.

The paleomagnetic data developed by Halls [1974] come from the Osler Volcanic Group in the Nipigon Straits region. The data were obtained from the uppermost part of the stratigraphic succession below the angular unconformity on Puff Island that separates the flows of reversed polarity from younger flows of normal polarity. The Halls [1974] data of reversed polarity (N = 25, http://earthref.org/MAGIC/9518) come from a portion of the stratigraphy that should correlate with the upper third of the sequence at Simpson Island. Given that the flows are dominantly thin pahoehoe lavas and that the Halls [1974] study area is ca. 30 km to the west, those data should be comprised of distinct individual cooling units from the Simpson Island flows. Watson and bootstrap tests for a common mean between the Halls [1974] data of reversed polarity and data from the upper third of the stratigraphy at Simpson Island are positive indicating that the populations of directions cannot be distinguished from one another—consistent with this stratigraphic correlation. In contrast, tests for a common mean between the reversed polarity Halls [1974] data and the lower third of the Simpson Island stratigraphy fail. This result indicates that the populations are statistically distinct,
building on the result that there was significant paleogeographic change recorded by the Osler Volcanic Group. Given that the data from the upper third of the stratigraphy on Simpson Island correlate stratigraphically and share a common mean with the Hall's [1974] data from the Nipigon Straits region, we can calculate a mean paleomagnetic pole for the upper portion of the reversed Osler Volcanic Group stratigraphy (42.5°N, 201.6°E, A95 = 3.7°, N = 59). This pole can be assigned an approximate age of 1105 ± 2 Ma using the date from the Agate Point rhyolite [Davis and Green, 1997]. These new data from 84 flows of Osler Volcanic Group from the early stage of the Midcontinent Rift bolster evidence from the succession at Mamainse Point that the decrease in inclination through the history of the rift is a progressive change rather than a stepwise change across reversals. The interpretation of a stepwise change of inclination across reversals has been used to argue for reversal asymmetry at the time that was proposed to result from significant deviation from an axial dipole geomagnetic field [Pesonen and
The observed progressive change in inclination is more consistent with the hypothesis that inclination decrease is a result of fast equatorward motion of Laurentia [Davis and Green, 1997; Swanson-Hysell et al., 2009]. The poles calculated from the stratigraphic groupings of the Osler Volcanic Group at Simpson Island fit the progression along the path resulting from the lowermost reversed polarity zone at Mamainse Point, Ontario. These data sets combined indicate that there was significant plate motion during the early magmatic stage of North American Midcontinent Rift development.

At present, the duration of Osler Volcanic Group volcanism remains poorly constrained given the uncertainties on the ages and the lack of a dated unit that can firmly be demonstrated to be extrusive near the base of the group. This reality makes it difficult to go from the plate motion inferred from the Osler Volcanic Group poles to a rate estimate based on that motion alone. An alternative approach is to use the new mean pole for the upper part of the reversed Osler stratigraphy in conjunction with a pole developed from normal magnetized volcanics of the main magmatic stage. The main magmatic stage pole we use here is calculated using data developed by Tauxe and Kodama [2009] from lava flows of the southwest limb of the North Shore Volcanic group between the dated 40th Ave icelandite and the Palisade rhyolite (35.8°N, 182.1°E, A95 = 3.1°, N = 47; see details in supporting online information). The pole’s age can be taken to be the mean of the two dated units that bracket it (1098.4 ± 1.9 Ma for the 40th Ave icelandite and 1096.6 ± 1.7 Ma for the Palisade rhyolite) [Davis and Green, 1997]. We take a Monte Carlo simulation approach to determine rate uncertainty between the pair of poles where a large number of random draws are taken for age from a Gaussian distribution and for pole position from a Fisherian distribution allows for the 95% confidence range of the rate estimate to be determined (Figure 4; code and further details are provided in the supporting materials). Using these poles, this approach yields an estimated rate of latitudinal change of 24.0 cm/yr with a 95% confidence range of 15.2–44.4 cm/yr (Figure 4). This analysis supports previous arguments that Laurentia’s rate of motion likely exceeded 20 cm/yr during the period of rift development [Davis and Green, 1997; Swanson-Hysell et al., 2009]. Note that what is being considered is latitudinal change and that the rate estimate is therefore a minimum. Future geochronology with higher precision has the potential to significantly reduce the uncertainty of such rate estimates. Regardless, given that latitudinal drift rates are typically well below 10 cm/yr throughout the Phanerozoic [Torsvik et al., 2012], this motion can be considered to be quite rapid. The rate likely exceeded that of India’s motion between ~70 and ~50 Ma when rates were between 13 and 18 cm/yr. This period of India’s motion represents the fastest well-constrained sustained motion of continental lithosphere in the Mesozoic and Cenozoic Eras [van Hinsbergen et al., 2011; Torsvik et al., 2012].

From the beginning of Midcontinent Rift extrusive volcanism in the early magmatic stage (ca. 1110–1105 Ma) to the voluminous volcanism associated with the main magmatic stage of rift development (ca. 1100–1095 Ma) there was ca. 25° of latitudinal motion of Laurentia. The latitudinal change implied
by the lower third Simpson Island Osler Volcanic Group pole to the North Shore Volcanic Group pole described above is 26.7° with the Osler Volcanic Group poles being on the first part of this polar wander trajectory during the early magmatic stage. The voluminous volcanism appears to have been concentrated in the Lake Superior region both during the early magmatic stage and during the eruption of the thick main stage volcanics in the central graben. Arguments for a plume origin for Midcontinent Rift volcanism have argued that a plume is necessary to explain isotopic signatures in lava flows (Nd isotopic data [Nicholson et al., 1997]; Re-Os isotopic data [Shirey, 1997]) and as a heat source for generating the large volumes of basaltic magma associated with the rift [Cannon, 1992]. If a long-lived plume was in a fixed position relative to Earth’s spin axis and did not become significantly diverted upon reaching the lithosphere, large relative motion of Laurentia could make it unable to continue to be a source of melt to the rift. One possibility is that Laurentia and a deep-seated mantle plume traveled in unison to lower latitudes as a result of large-scale rapid true polar wander. The motion implied by Keweenawan Track poles has been interpreted as a result of true polar wander [Evans, 2003; Mitchell et al., 2012]. To reconcile an interpretation of rapid plate motion through differential plate tectonic motion, rather than true polar wander with a continued plume contribution, Davis and Green [1997] proposed that a plume head drifted with continental lithosphere. This idea can be considered in the context of the model of “upside-down drainage” for positively buoyant plume material wherein relief at the base of the lithosphere directs lateral flow [Sleep, 1997; Ebinger and Sleep, 1998]. Given the significant lithospheric thinning associated with rifting, plume material could continue to be directed into the rift sustaining magmatism despite ongoing plate motion. Such lateral flow could also explain the widespread coeval volcanism recorded by mafic intrusions in SW Laurentia [Heaman and Grotzinger, 1992; Weil et al., 2003]. Whether continued lateral flow into the rift is feasible with this scale of differential plate motion depends on the paleogeographic configuration at the time and whether continuous continental lithosphere extended beyond cratonic Laurentia as part of a nascent Rodinia supercontinent.

At the same time that the Midcontinent Rift was initiating in Laurentia the following igneous provinces were emplaced on four other cratons:

1. The Umkondo large igneous province of the Kalahari Craton (many isotope dilution thermal ionization mass spectrometry (ID-TIMS) U-Pb dates on zircon and baddeleyite between 1112 and 1108 Ma) [Hanson et al., 2004].
2. Thick and extensive gabbro-norite (GN) dikes exposed in the southwest Angola portion of the Congo Craton (one of which has an ID-TIMS U-Pb date on baddeleyite of 1110.3 ± 2.5 Ma) [Ernst et al., 2013].
3. The Mahoba suite of dikes of the India Craton (one of which, the “Great Dike of Mahoba” has an laser ablation U-Pb date on zircon of 1112.7 ± 7.4 Ma; Pradhan et al. [2012]).
4. A putative ca. 1110 Ma large igneous province in the southwest portion of the Amazonia craton inferred from dates on two intrusions (ID-TIMS U-Pb date on baddeleyite from the Rincón del Tigre intrusion of 1110.4 ± 1.8 Ma and a sill within Aiguepe sediments with a ID-TIMS U-Pb date on baddeleyite of 1111.5 ± 1.5 Ma [Hamilton et al., 2012]).

This contemporaneous voluminous volcanism on five late Mesoproterozoic cratons is coincident with the onset of Laurentia’s rapid plate motion. This temporal correlation suggests that this time period was characterized by particularly vigorous plume activity that in addition to driving large igneous province development may be connected to the driving forces that resulted in rapid differential plate motion and/or true polar wander.

5. Conclusion
Lava flows of the Osler Volcanic Group below the Puff Island unconformity erupted during the early magmatic stage of Midcontinent Rift development during a time interval characterized by reversed magnetic polarity. New paleomagnetic data from 84 Osler Volcanic Group lava flows reveal a significant decrease in paleomagnetic inclination through the sequence of lava flows. These results support the hypothesis that the difference between the steep paleomagnetic inclinations characteristic of the early magmatic stage and the relatively shallower inclinations of the main magmatic stage throughout the rift are the result of rapid plate motion that was progressively recorded by the Osler Volcanic Group.
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References


